

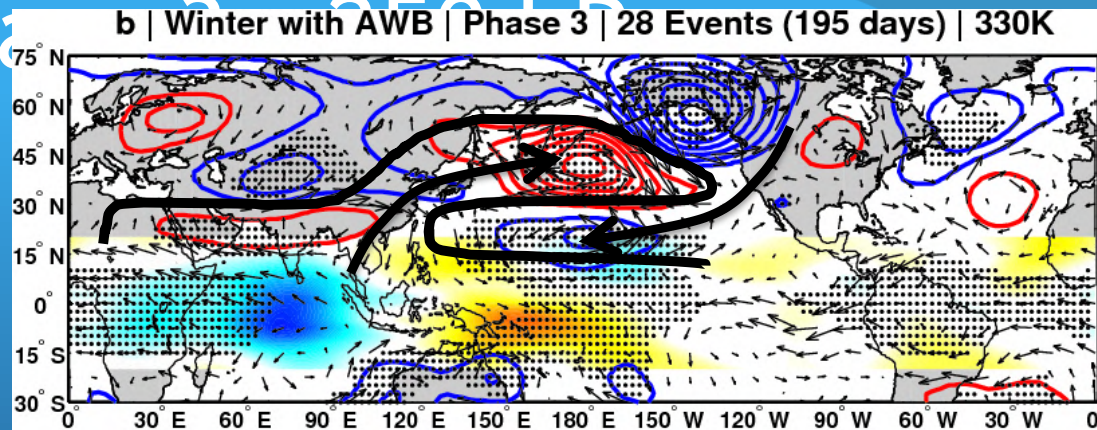
The MJO and Extratropical Wave Breaking

- Focus Northern Winter
- Identify continuously active events from the RMM index from phase 2 forward > Amp 0.5
- Identify those events that include central Pacific wave breaking phase 3
- Compare progression of events with wave breaking to events without
 - MacRitchie and Roundy (2016)

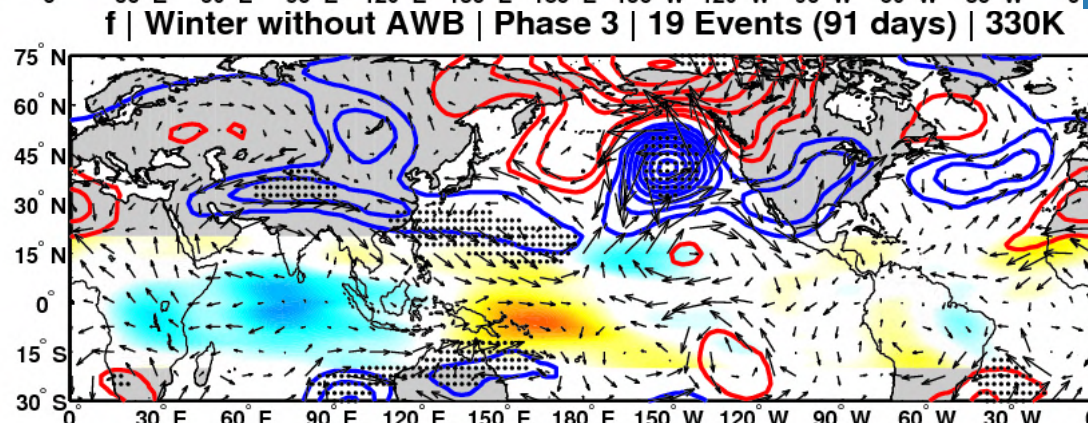
From MacRitchie and Roundy (2016, QJRMS)

Phase 3

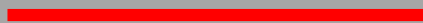
RMM Amp:
1.85



RMM Amp:
1.27



Positive Height Anomalies (15 m):



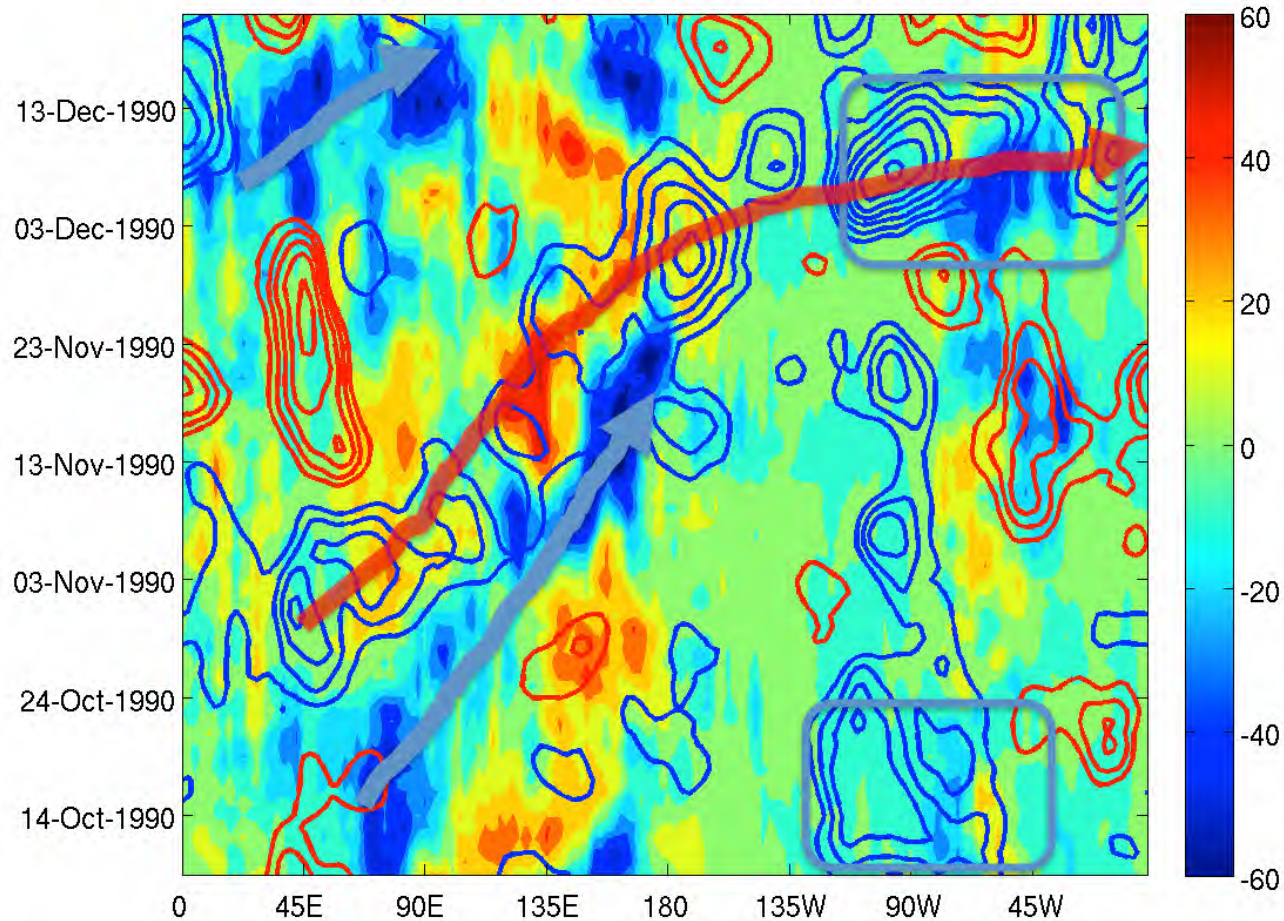
Negative Height Anomalies (15 m):



OLR Anomalies:



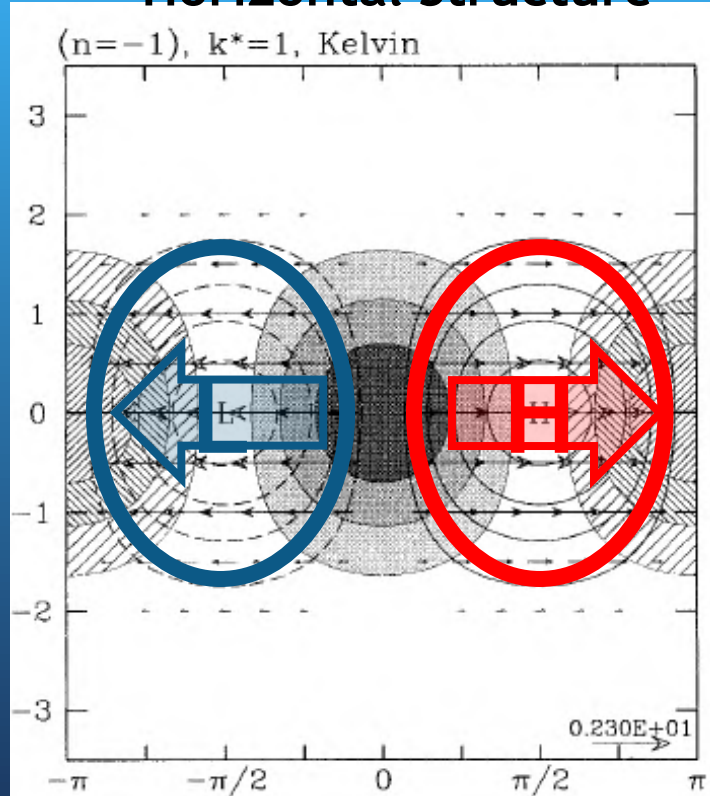
Equatorial OLR and 200 hPa Zonal Wind Anomalies



Roundy 2015

Theoretical Kelvin Wave Structure

Horizontal Structure



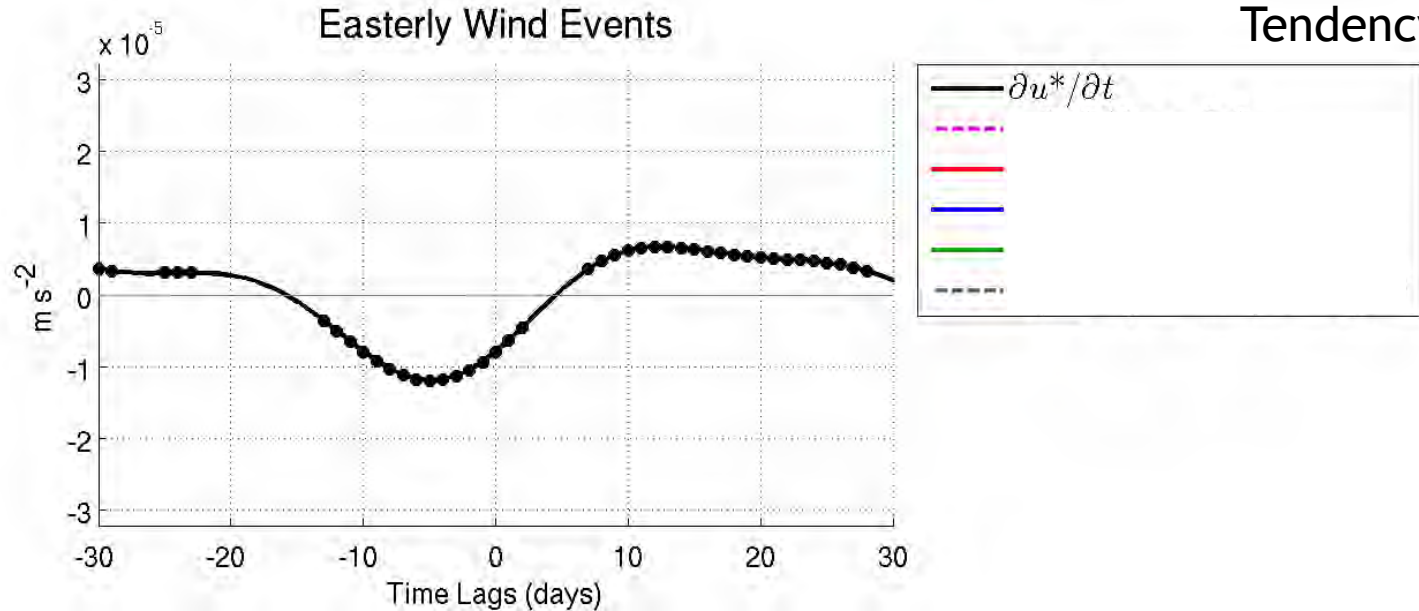
$$\frac{\partial u}{\partial t} = -g \frac{\partial \Phi}{\partial x}$$

: Zonal Momentum Equation

Circumnavigating MJO Circulation: Zonal Momentum Budget at 200 hPa

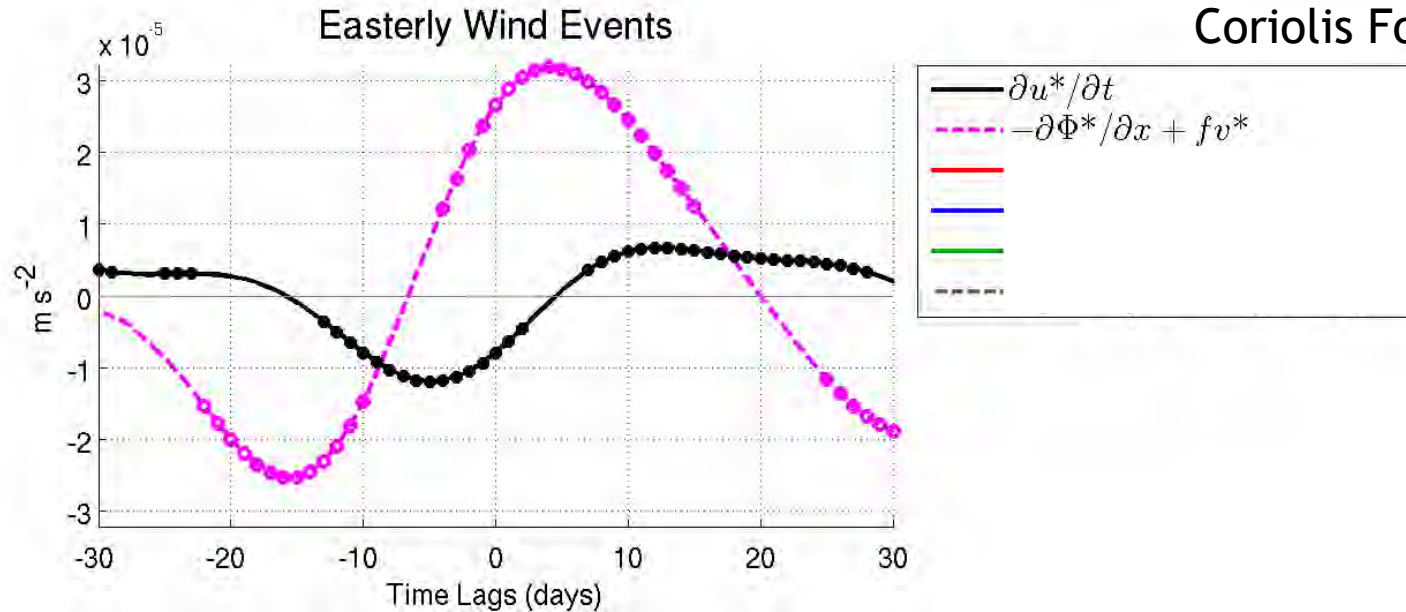
$$\frac{\partial u}{\partial t} = -\vec{v} \cdot \nabla u - g \frac{\partial \Phi}{\partial x} + f v + X$$

Intraseasonal Zonal Wind Tendency



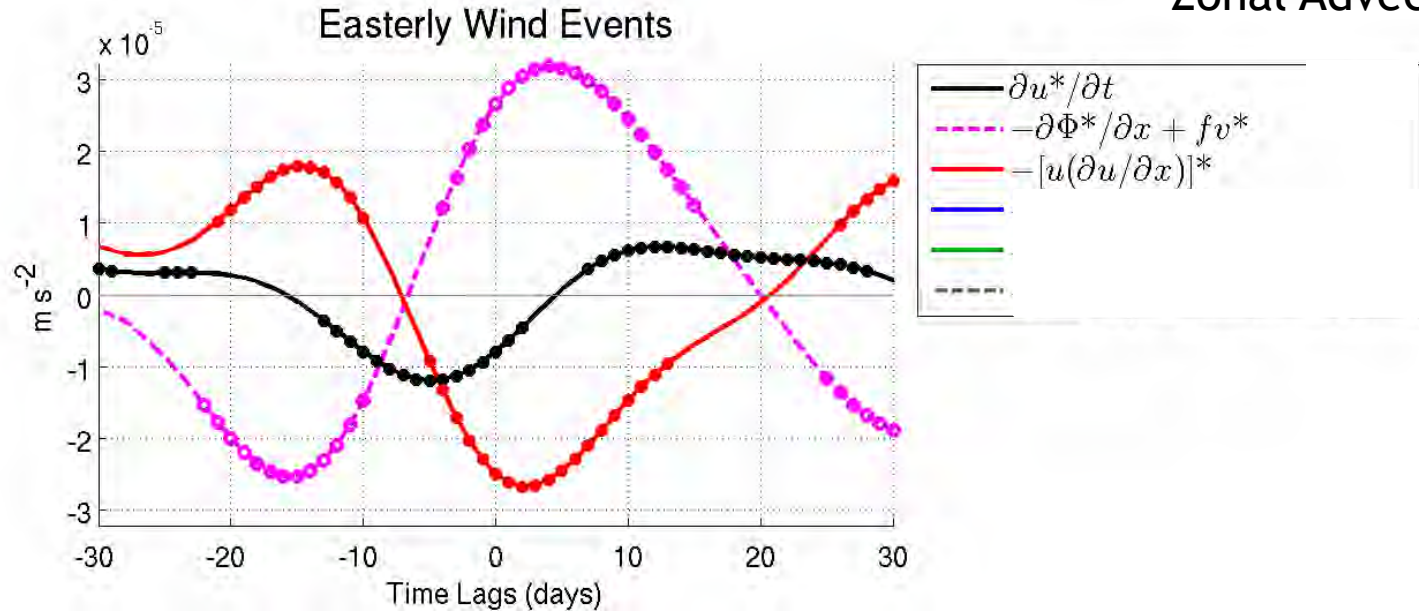
10°S-5°N, 110°W-80°W Averaged u Budget Terms

Pressure Gradient Force and
Coriolis Force



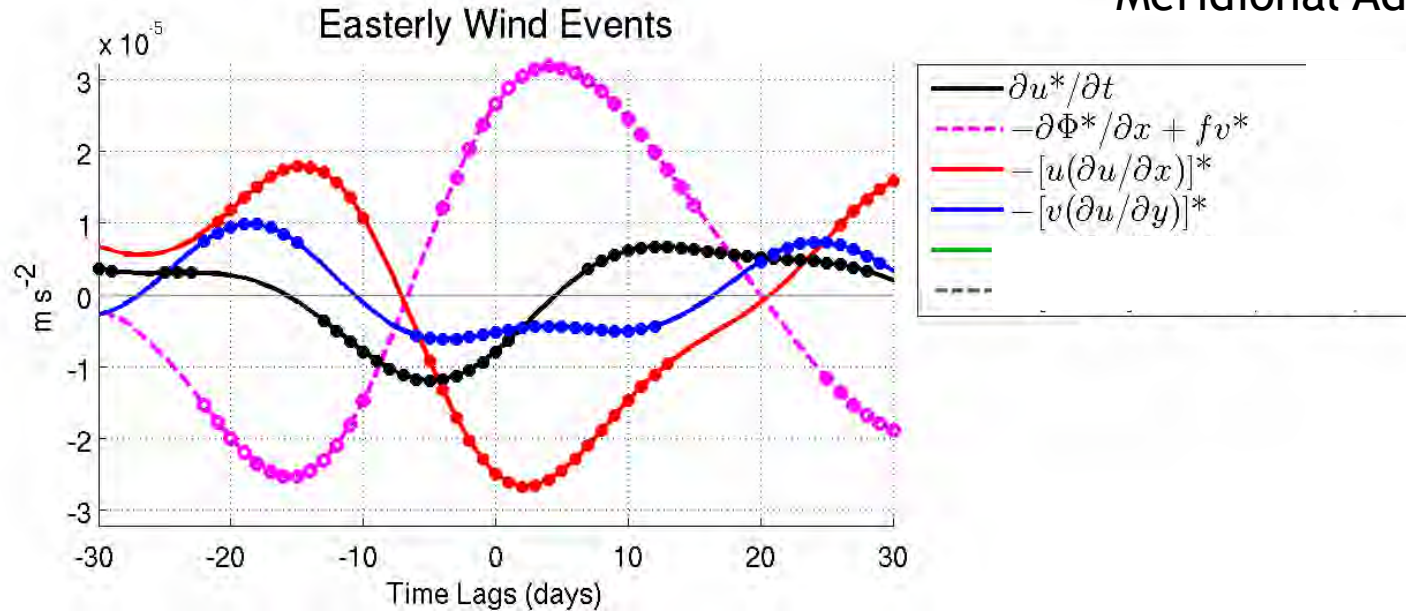
10°S-5°N, 110°W-80°W Averaged u Budget Terms

Zonal Advection



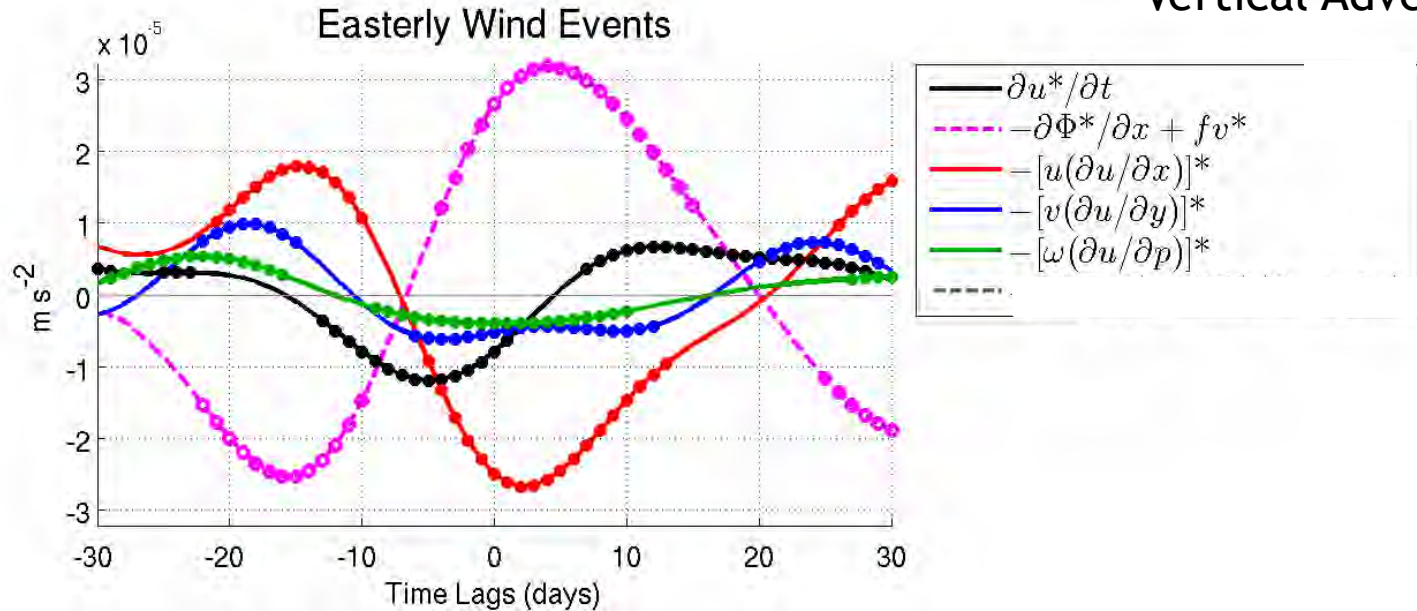
10°S-5°N, 110°W-80°W Averaged u Budget Terms

Meridional Advection



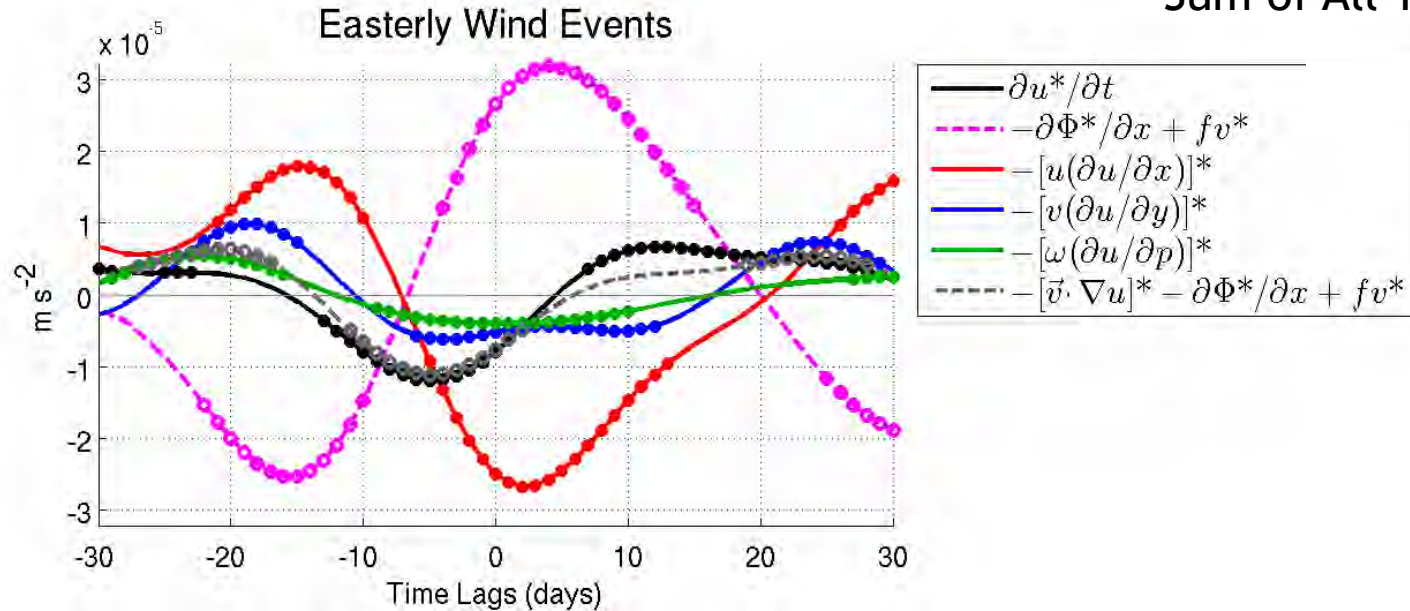
10°S-5°N, 110°W-80°W Averaged u Budget Terms

Vertical Advection



10°S-5°N, 110°W-80°W Averaged u Budget Terms

Sum of All Terms



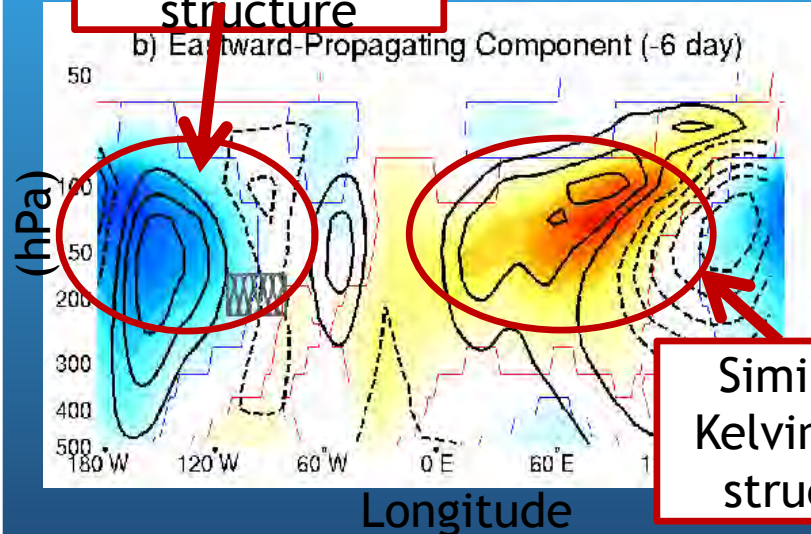
Sakaeda and Roundy (2015, 2016a,b,c)

10°S-10°N Averaged Longitude-Pressure Cross Sections

Day -6

NOT Similar to
Kelvin wave
structure

Pressure
(hPa)

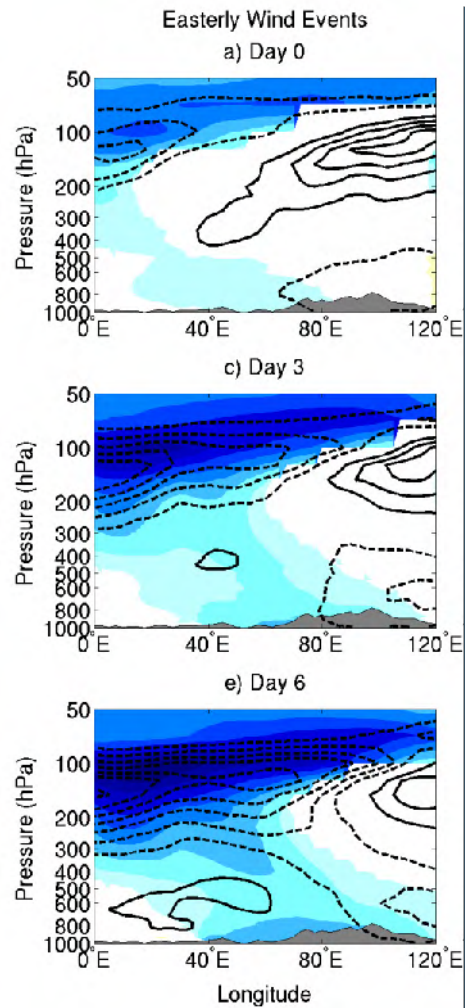


Sakaeda and Roundy (2015)

Similar to
Kelvin wave
structure

Shading: intraseasonal zonal wind anomaly

Black contour: intraseasonal geopotential height anomaly at 2 m interval



Sakaeda and Roundy (2015)

Interaction with the extratropics may be associated with

- Changes in amplitude and spatial distribution of MJO convection
- Changes in the phase speed of MJO convection and associated global circulation anomalies

Three evidence pathways:

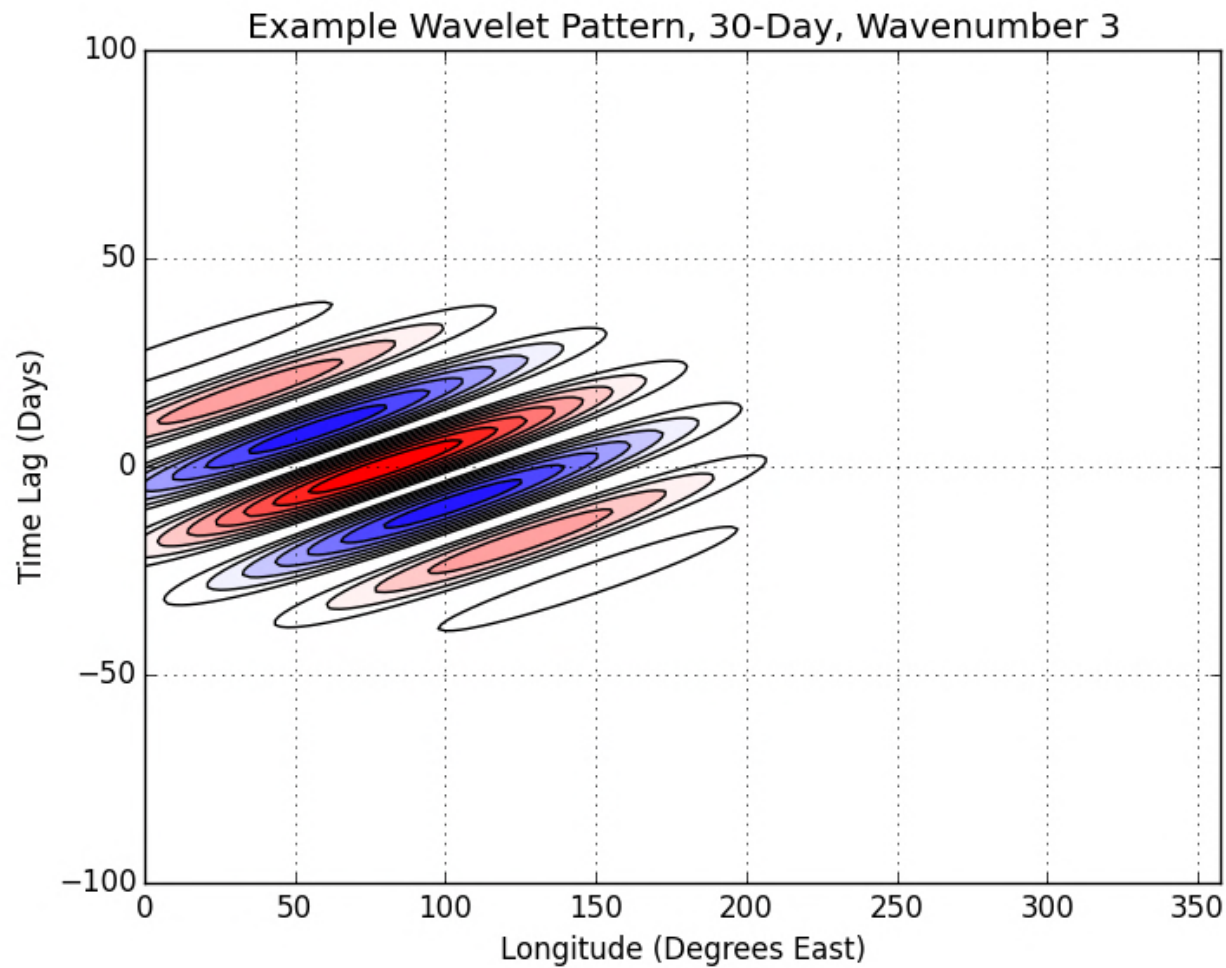
- Regress fields of data against OLR data filtered by wavelet analysis for specific phase speeds,
- Modified regression against the RMM index, and
- Simple case analysis

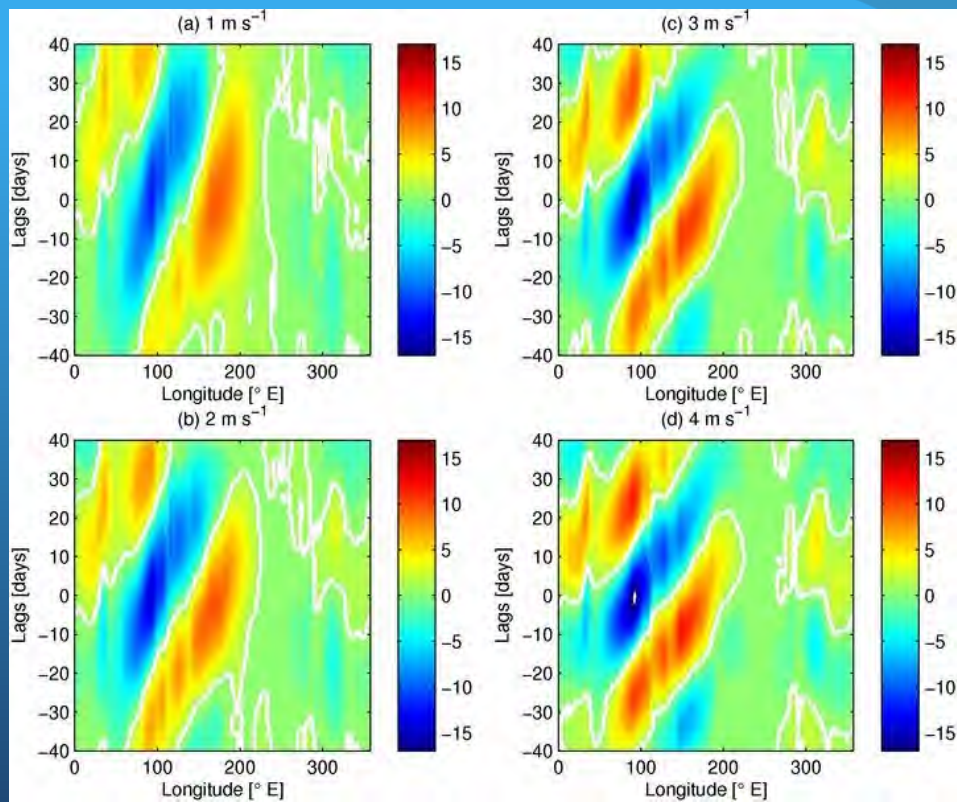
Space Time Wavelet Analysis

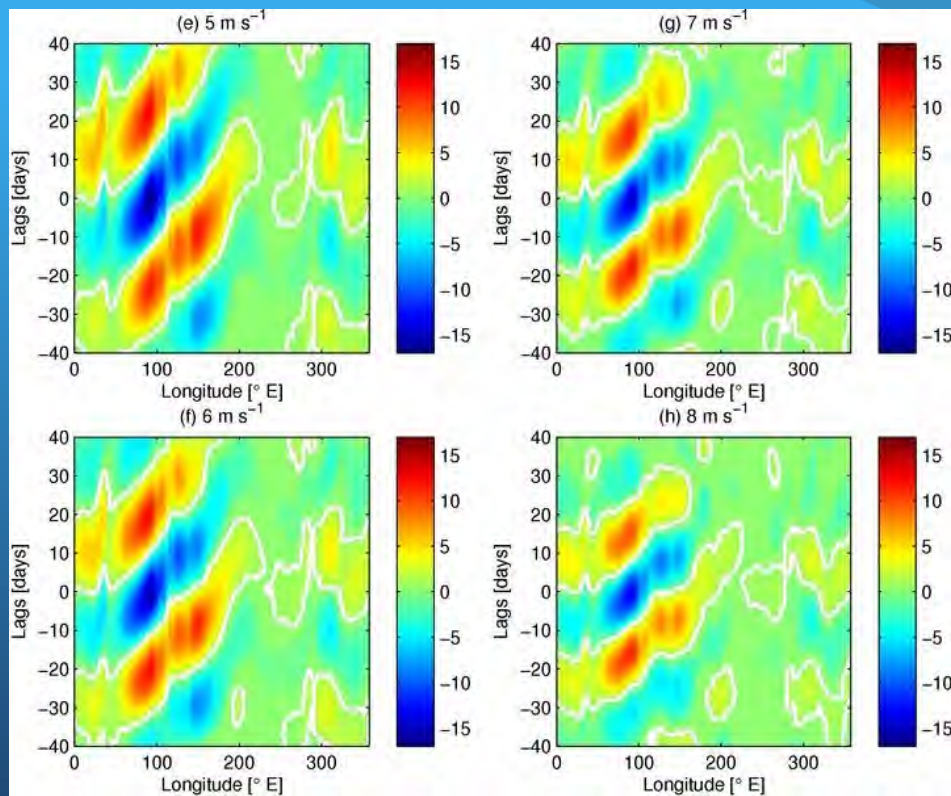
$$\psi = \frac{1}{\gamma} \cos(2\pi(kx - F_c t)) \exp\left(\frac{-x^2}{F_{Bx}} - \frac{-t^2}{F_{Bt}}\right) \quad (1), \text{ where}$$

$$\gamma = \sum_{x=0}^{144} \sum_{t=-100}^{100} \exp\left(\frac{-x^2}{F_{Bx}} + \frac{-t^2}{F_{Bt}}\right)$$

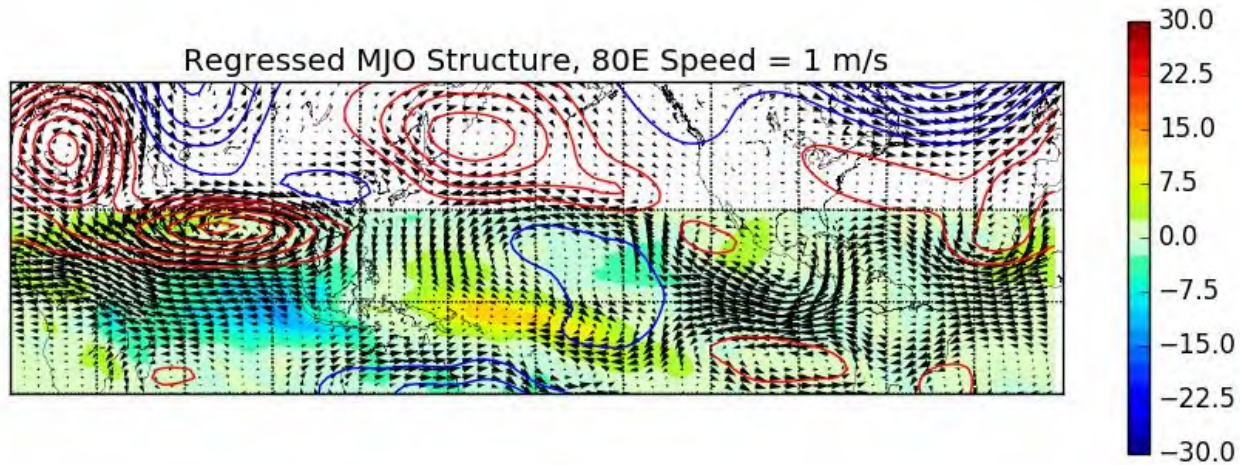
Thanks to grad students Bob Setzenfand and Kaitlyn Krzyzaniak



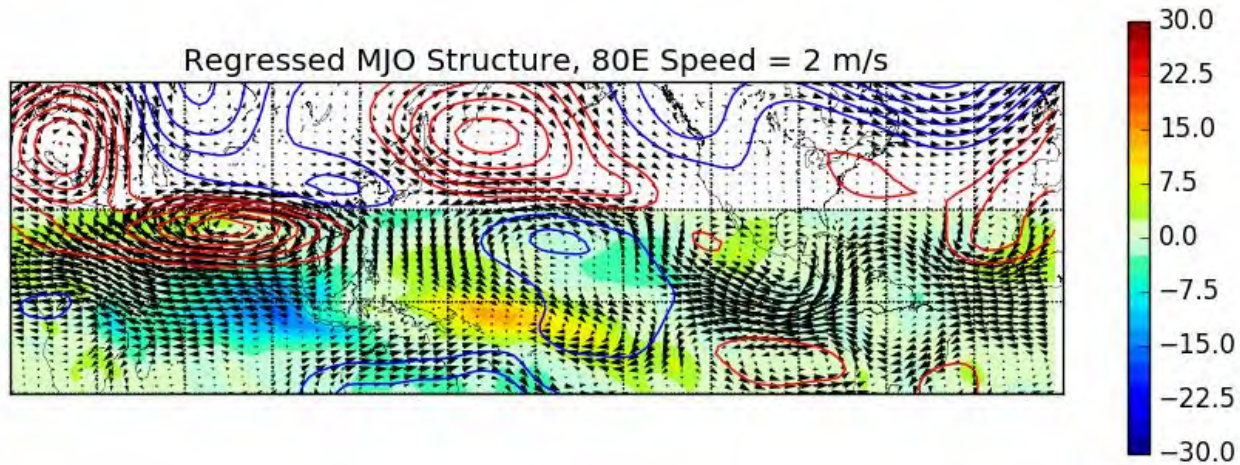




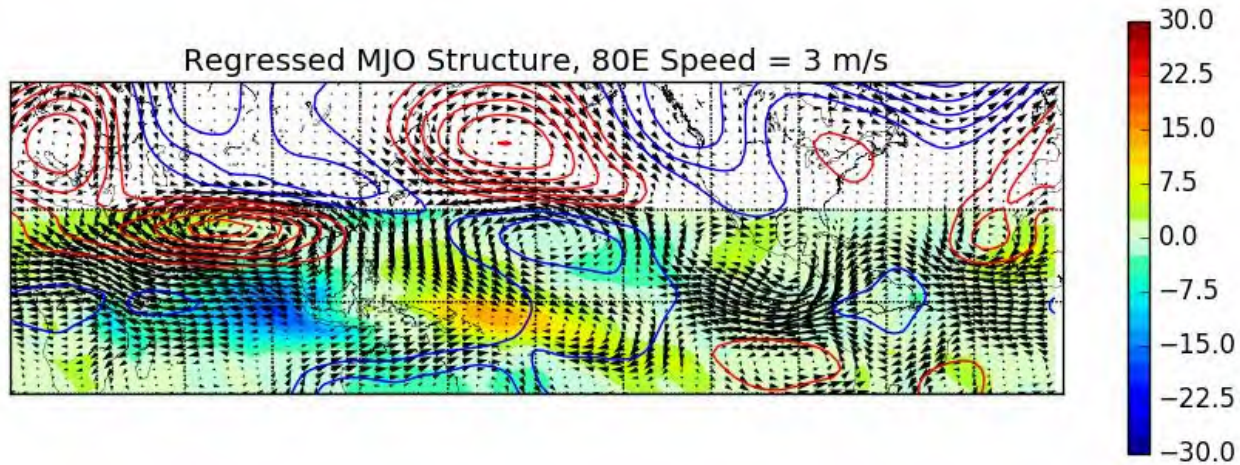
200 hPa Height Anomaly



200 hPa Height Anomaly

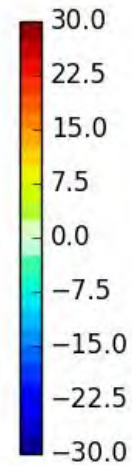
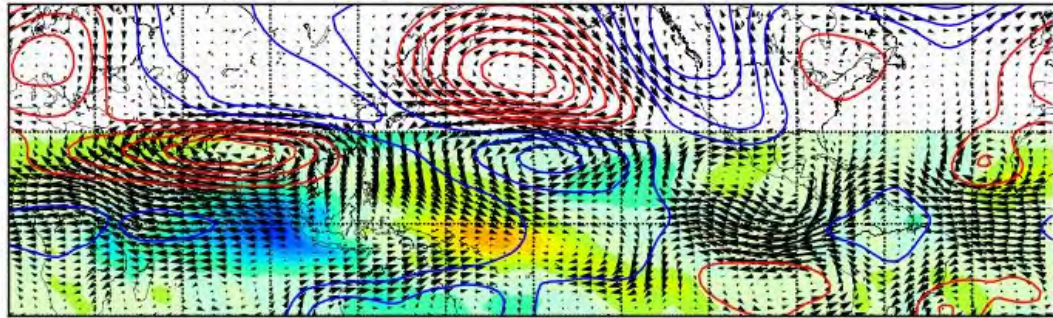


200 hPa Height Anomaly

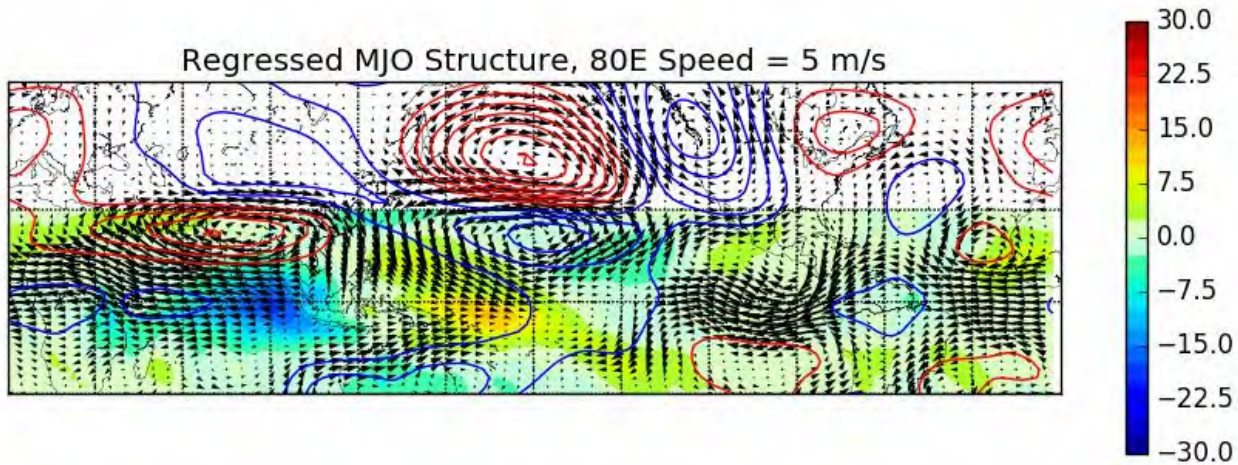


200 hPa Height Anomaly

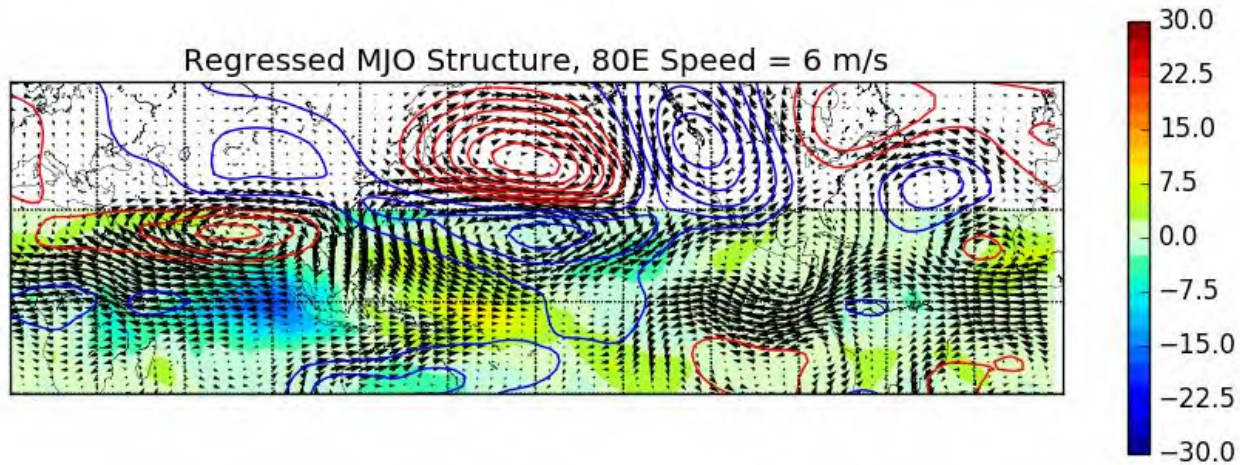
Regressed MJO Structure, 80E Speed = 4 m/s



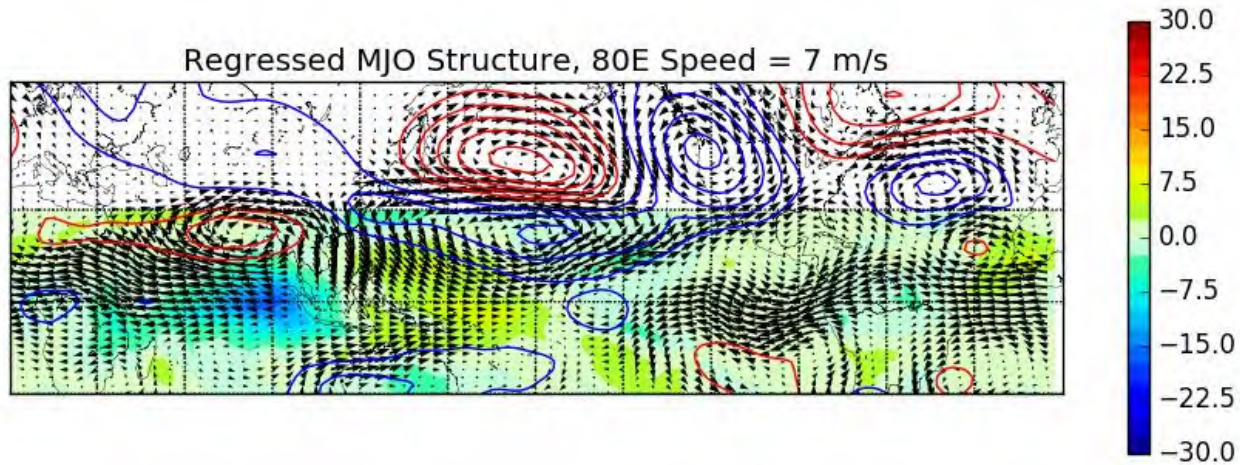
200 hPa Height Anomaly



200 hPa Height Anomaly



200 hPa Height Anomaly

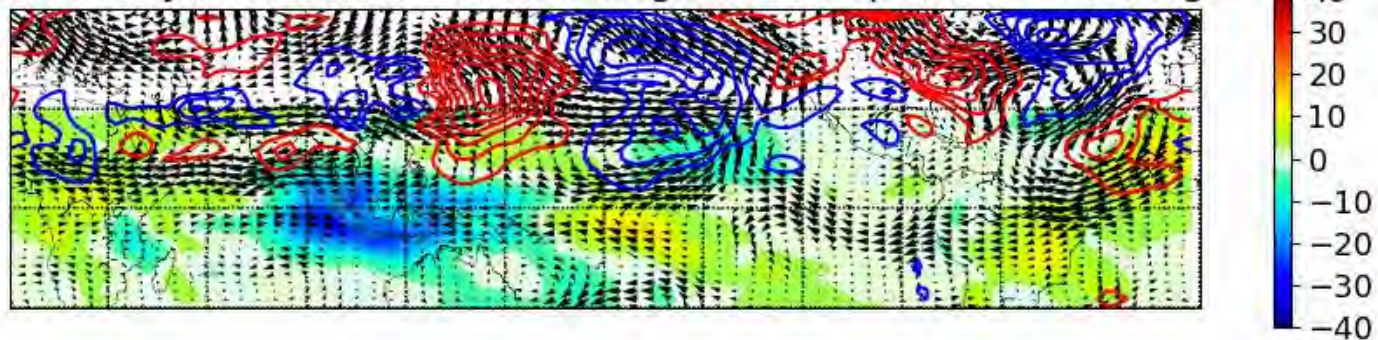


Supports conclusions of
Yadav and Straus (2017, MWR)

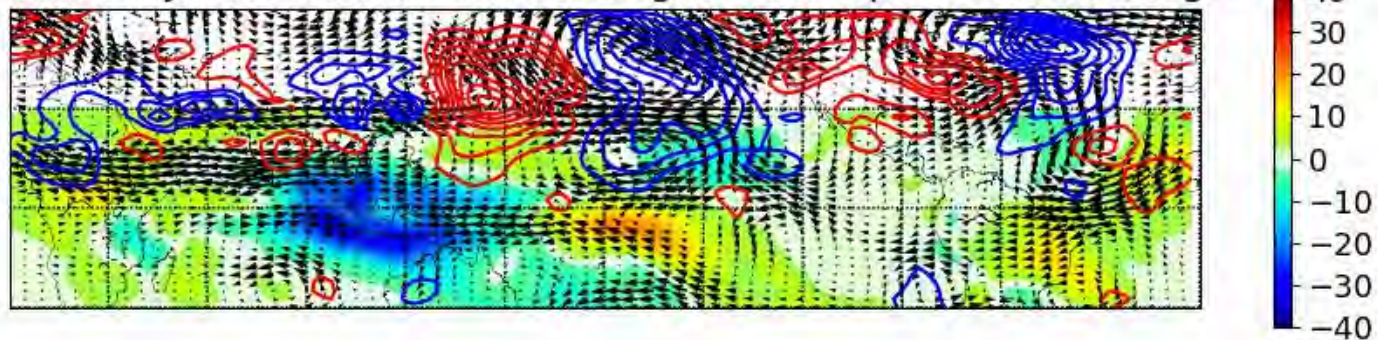
Wavelet Regression of Horizontal Advection Terms

- Regress data for horizontal advection of zonal wind against wavelet filtered MJO index at select phase speeds
- Compare magnitude of regressed anomalies with the selected phase speed

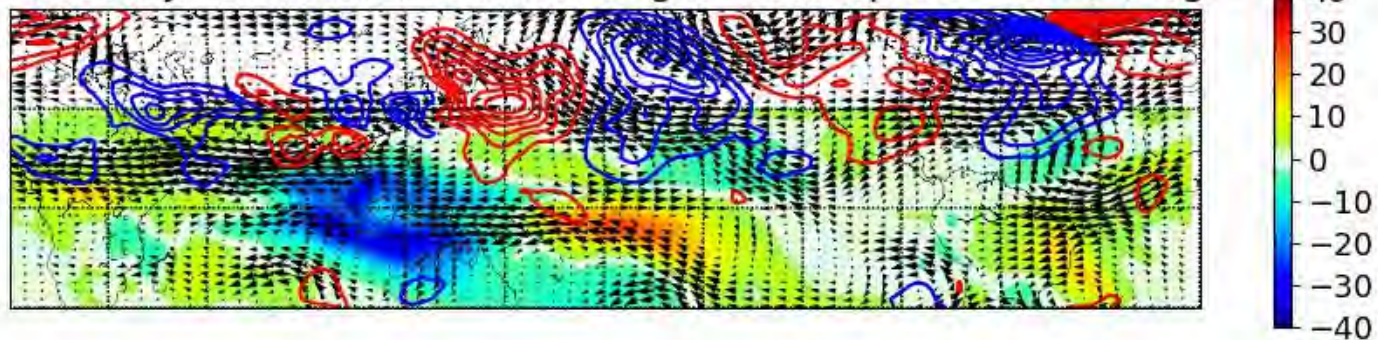
Regressed MJO Structure, 400 hPa Height 110E Speed = 7 m/s, lag = 0



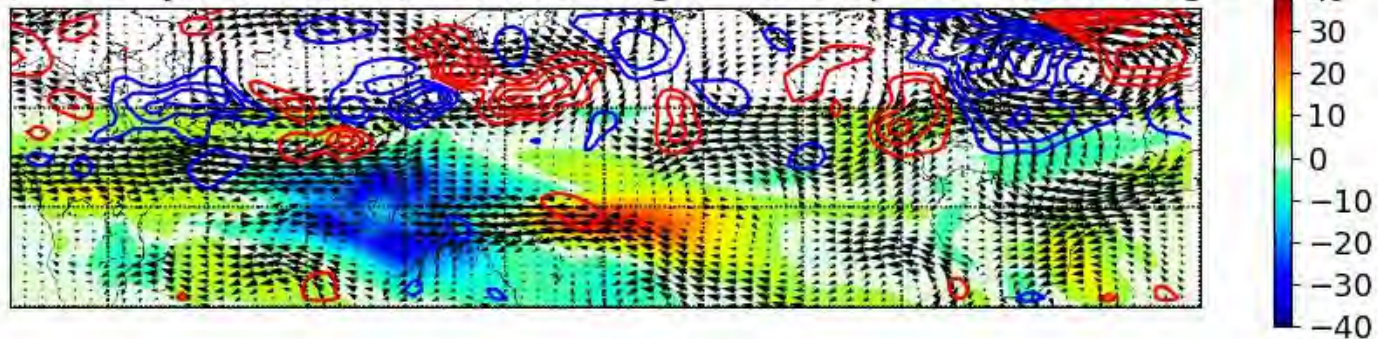
Regressed MJO Structure, 400 hPa Height 110E Speed = 5 m/s, lag = 0



Regressed MJO Structure, 400 hPa Height 110E Speed = 3 m/s, lag = 0



Regressed MJO Structure, 400 hPa Height 110E Speed = 1 m/s, lag = 0



Modified linear regression against RMM index in different base states

- Algorithm similar to method of seasonally varying regression slope coefficients by Roundy (2017 QJRMS)
- Regress December through February OLR and height anomalies against RMM index for a selected RMM phase during low or high background wind at highlighted location, and check the phase speed

- Slope coefficients are the ratio of covariance between predictor RMM index and predictand to variance in the predictor
- Model the covariance and variance based on external modulating factors, on this case, low frequency background winds
- Create the slope coefficients for different background conditions in the modulating factors

$$\hat{m}_p = \text{inv}(\mathbf{x}_p^T \mathbf{x}_p) \mathbf{x}_p^T \mathbf{y}_p$$

$$\mathbf{x}_{var} = \mathbf{x}_p * \mathbf{x}_p = \mathbf{x}_p^2.$$

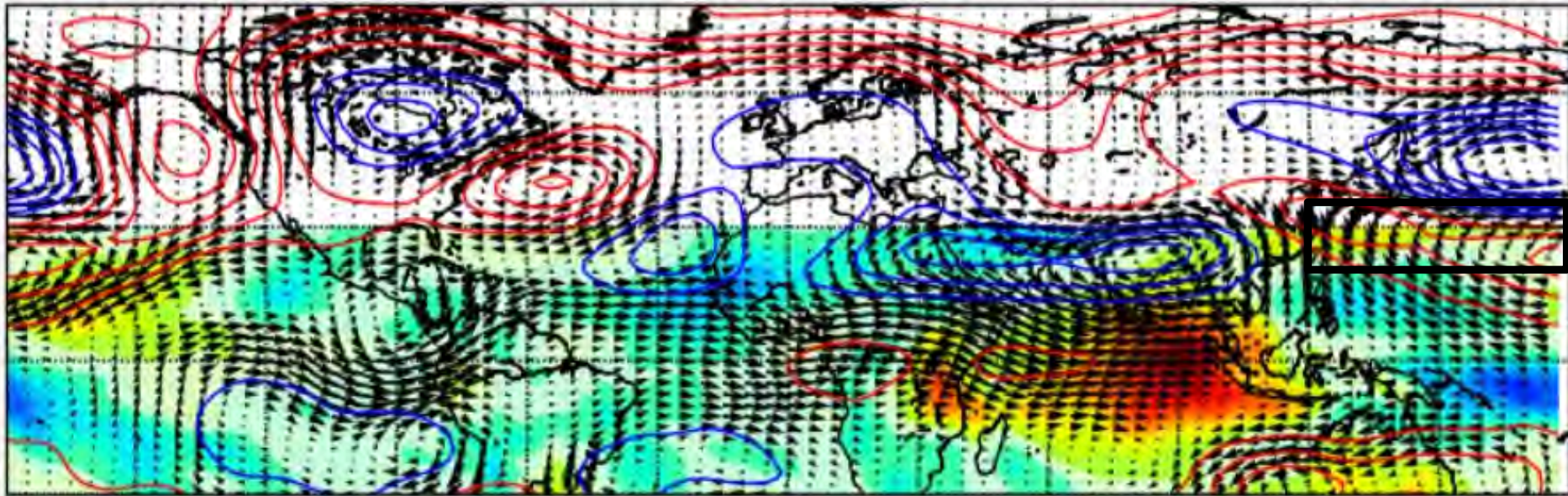
$$\hat{\mathbf{c}}_{var} = \text{inv}(\mathbf{X}^T \mathbf{X}) \mathbf{X}^T \mathbf{x}_{var}$$

$$\text{Variance} = \mathbf{X} \hat{\mathbf{c}}_{var}$$

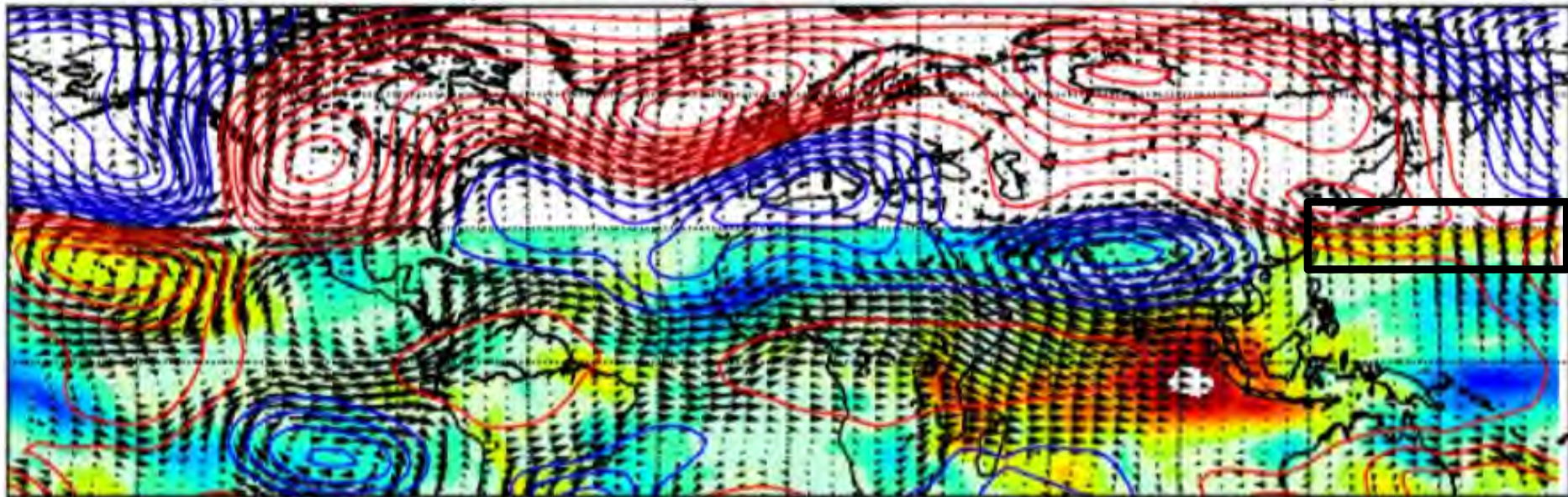
$$\hat{m}_p = \frac{\text{covariance}}{\text{variance}}$$

Roundy 2017, QJRM

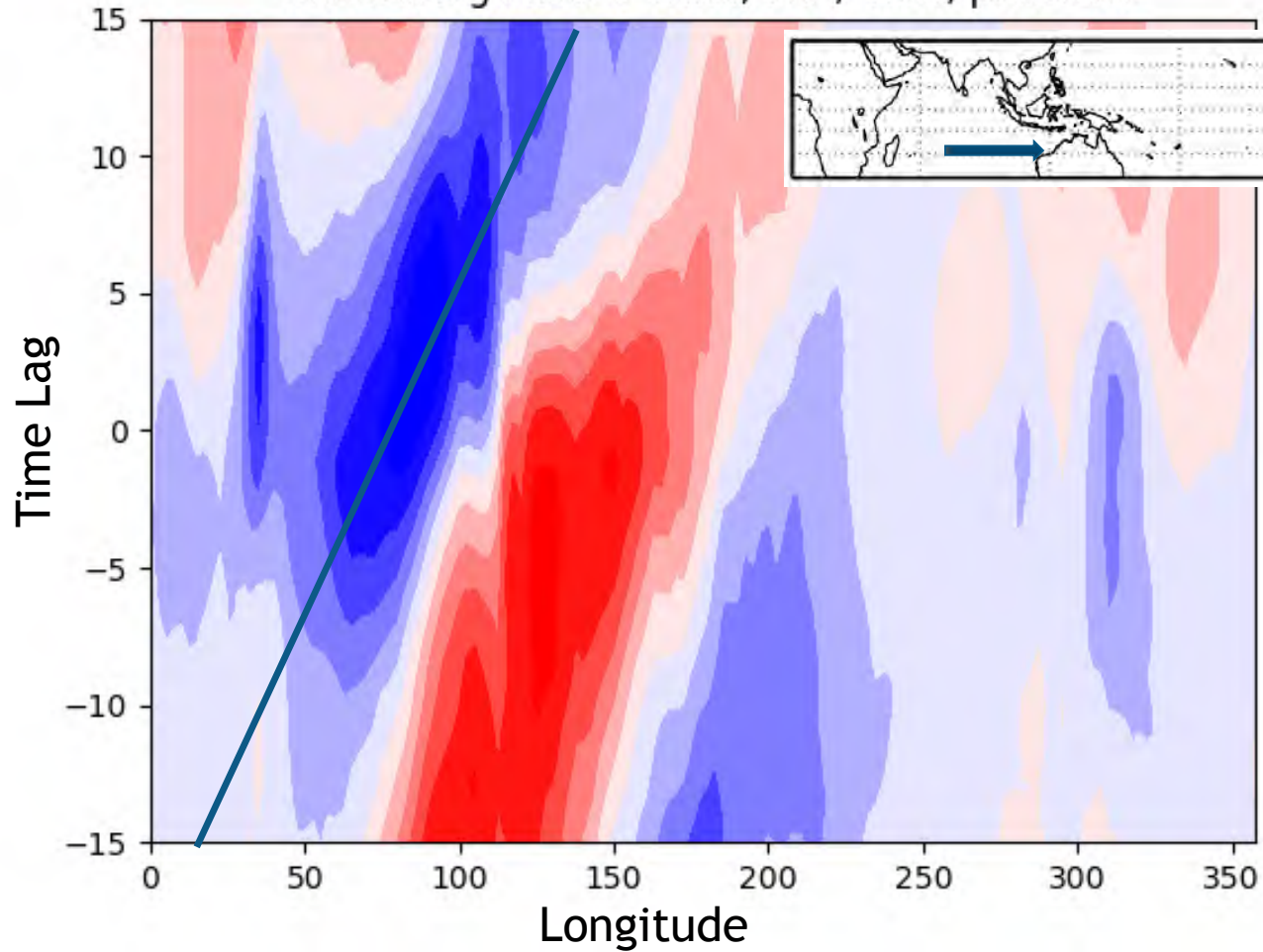
b. RMM Regression Map 200 Height (2 SDs) Phase 6 lead 5 days, 20 m/s



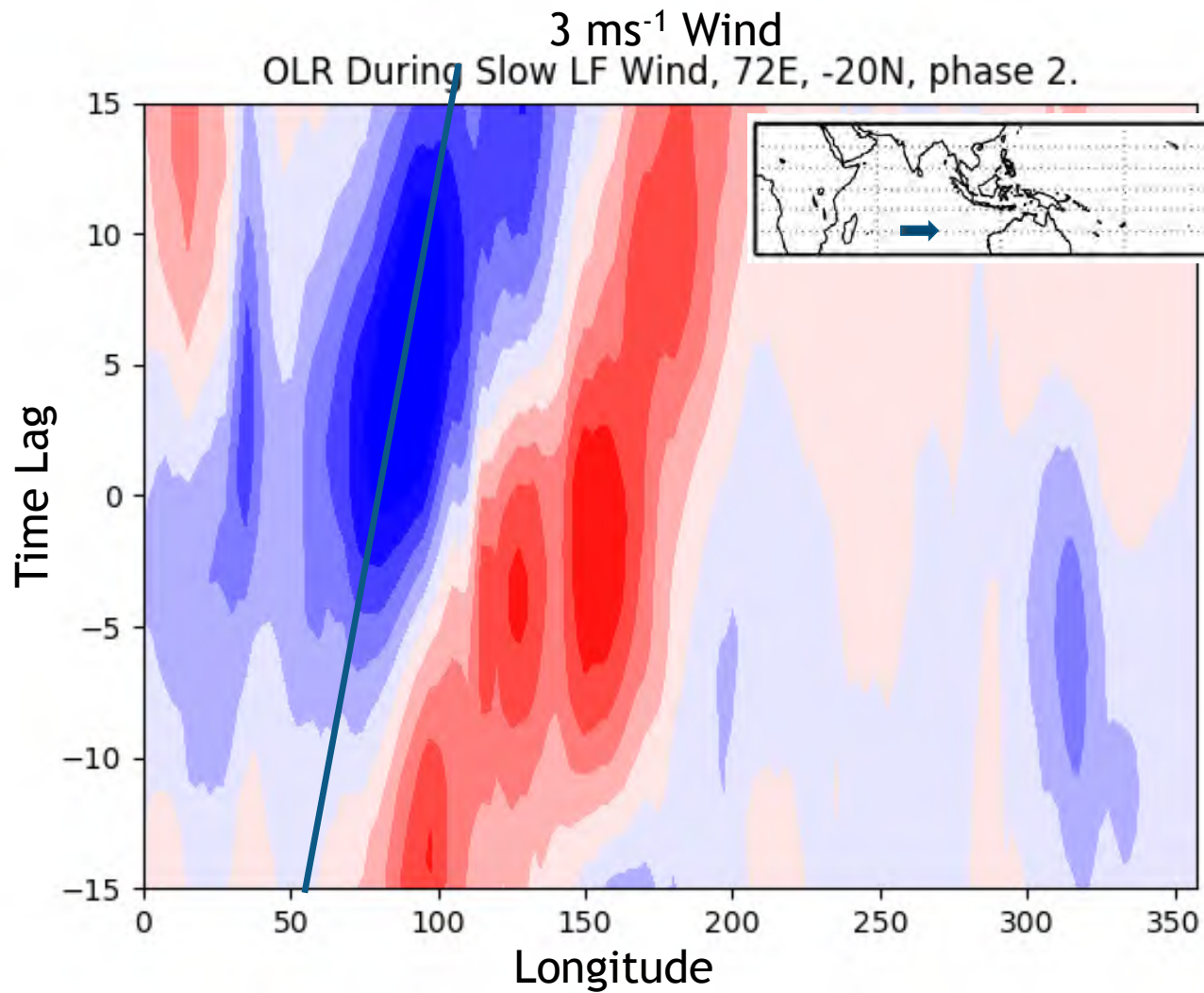
a. RMM Regression Map 200 Height (2 SDs) Phase 6 lead 5 days, 40 m/s



15 ms^{-1} Wind
OLR During Fast LF Wind, 72E, -20N, phase 2.

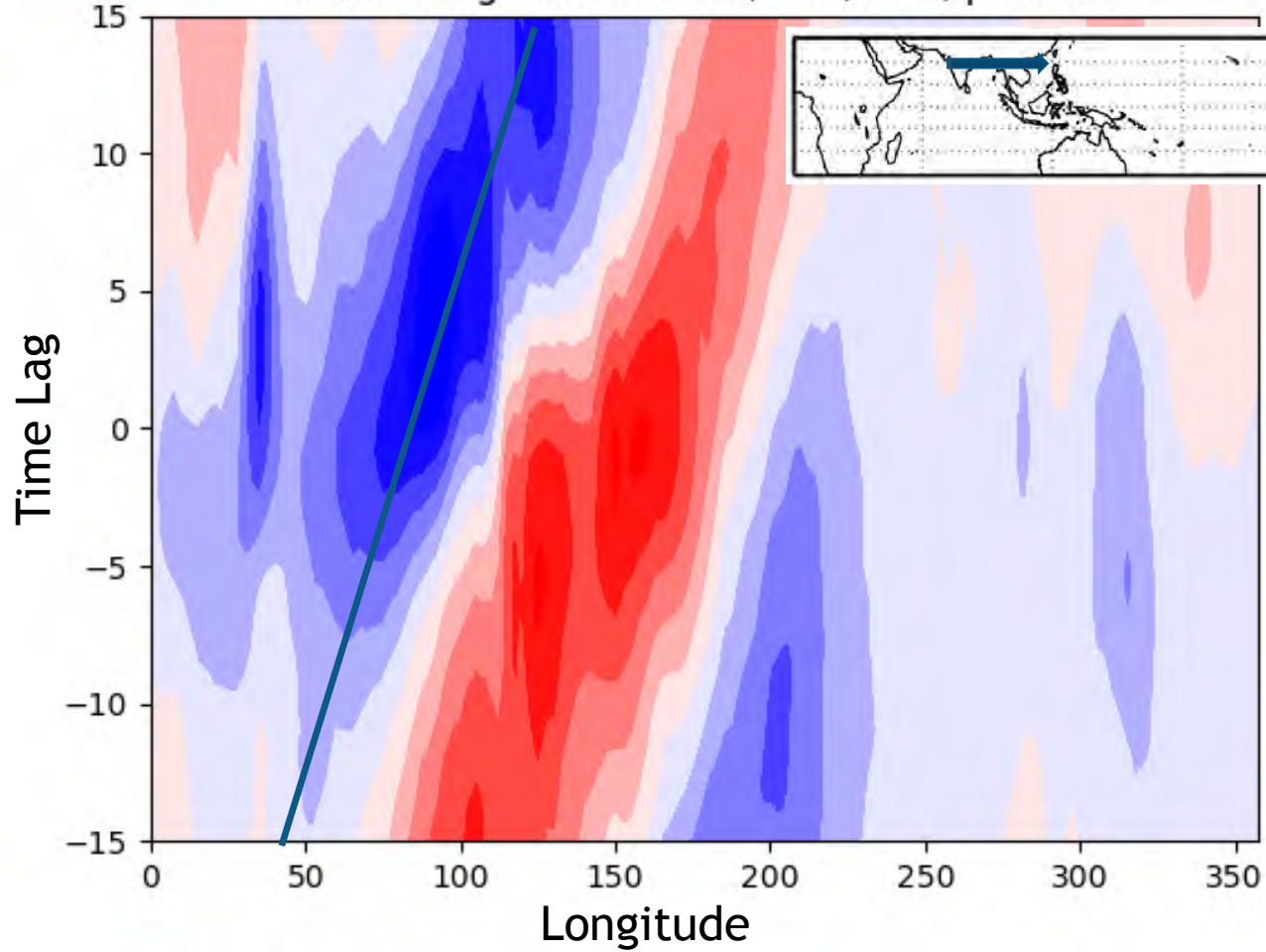


Phase
5 ms^{-1}

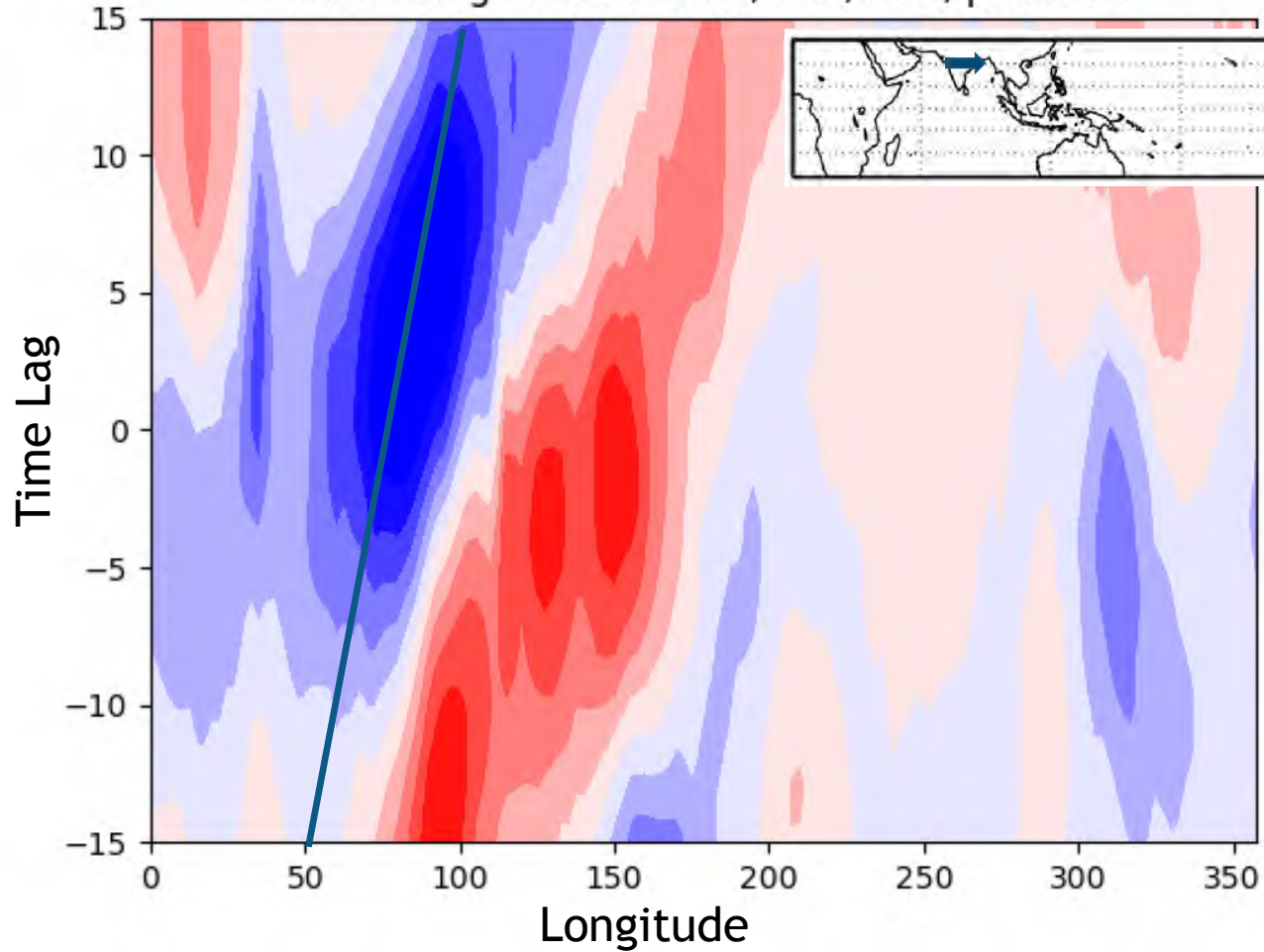


Phase
2ms⁻¹

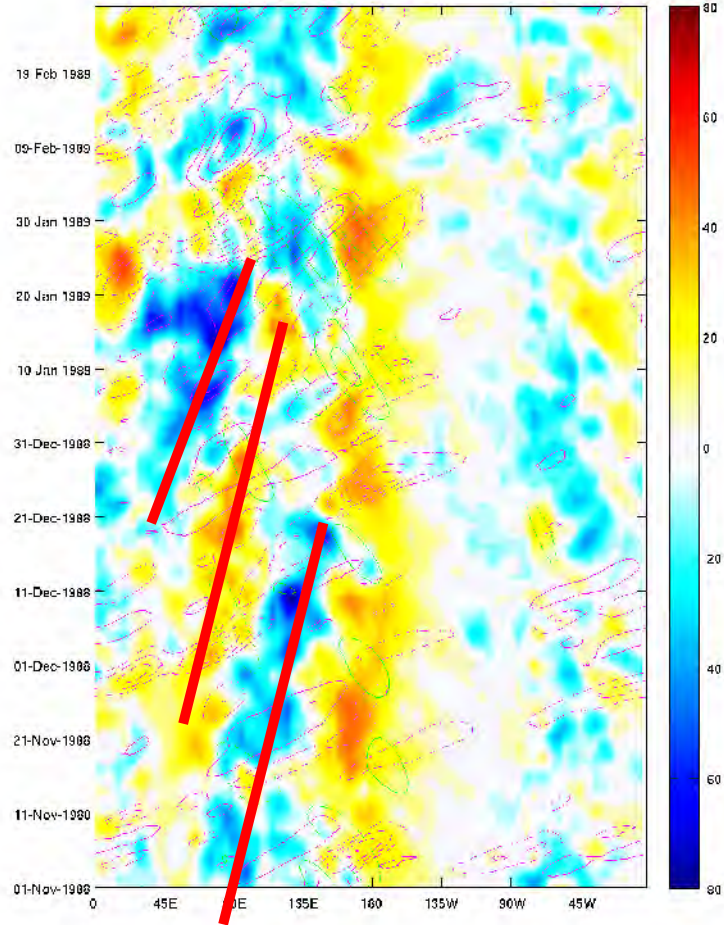
33 ms⁻¹ Wind
OLR During Fast LF Wind, 72E, 20N, phase 2.



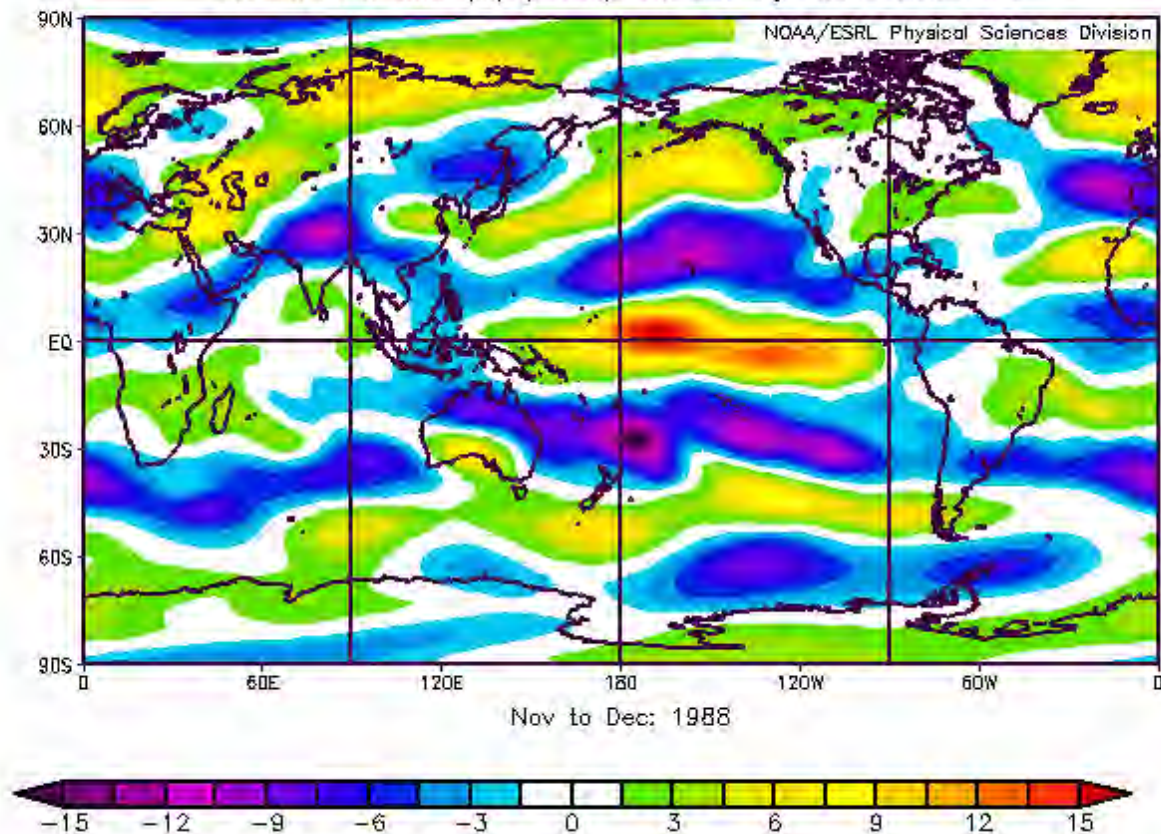
20 ms^{-1} Wind
OLR During Slow LF Wind, 72E, 20N, phase 2.



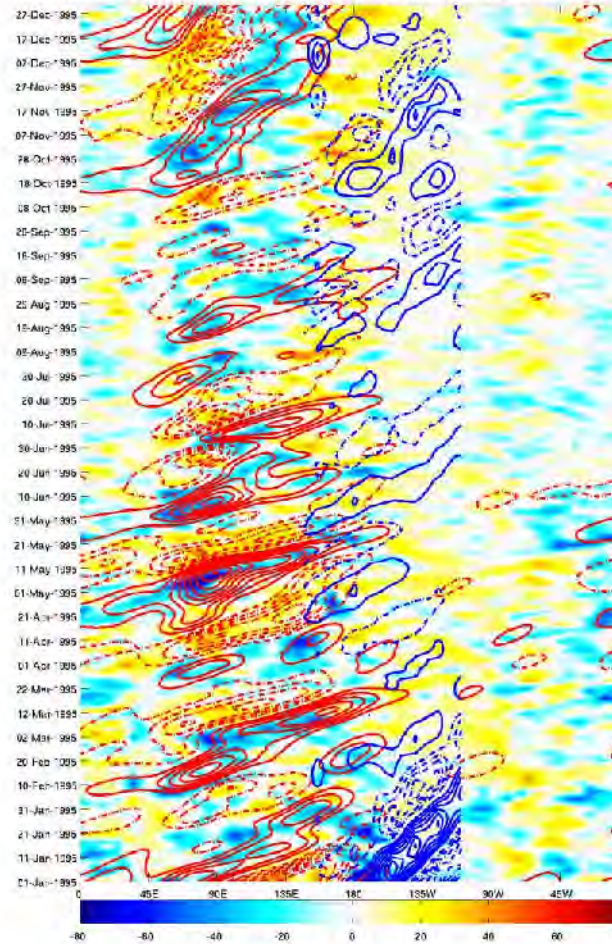
Unfiltered and Projected OLR Anomalies 7.5S to 7.5N



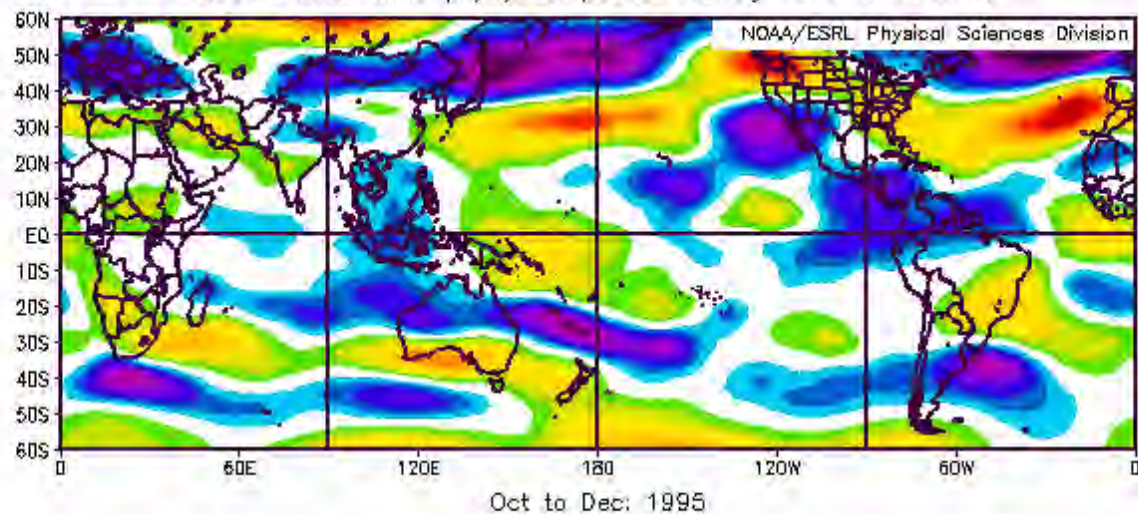
NCEP/NCAR Reanalysis
200mb Zonal Wind (m/s) Composite Anomaly 1981–2010 clima



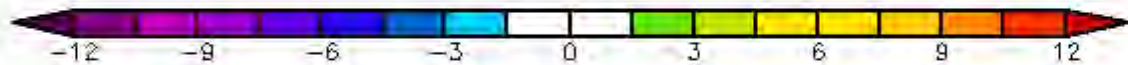
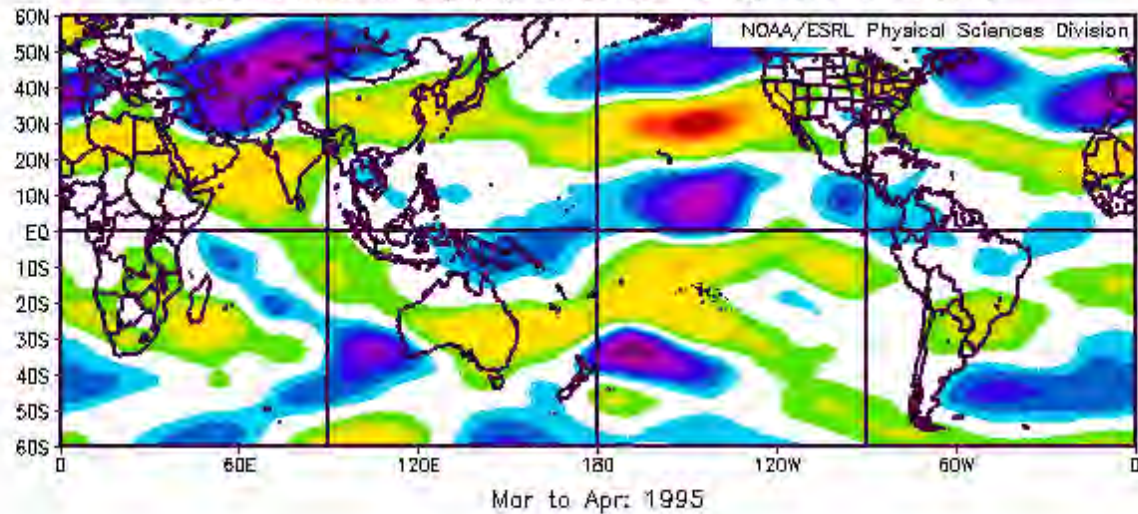
Unfiltered and Projected OLR Anomalies 7.5S to 7.5N



NCEP/NCAR Reanalysis
200mb Zonal Wind (m/s) Composite Anomaly 1981–2010 clima



NCEP/NCAR Reanalysis
200mb Zonal Wind (m/s) Composite Anomaly 1981–2010 clima

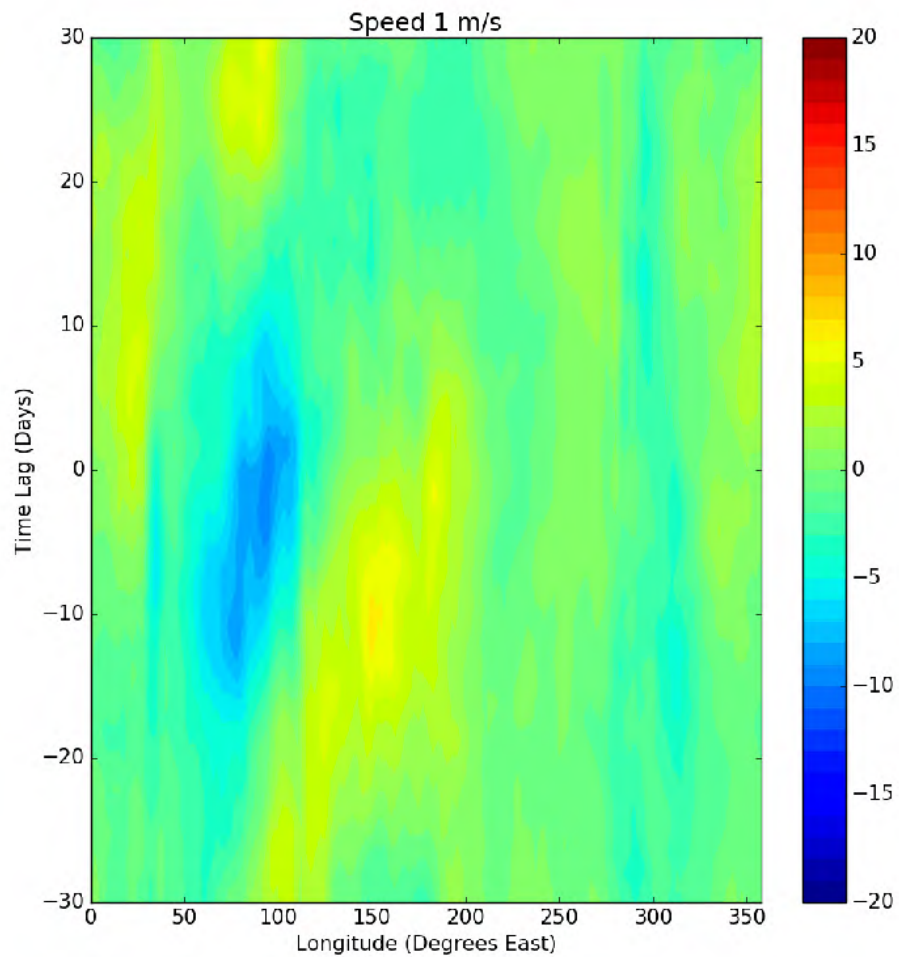
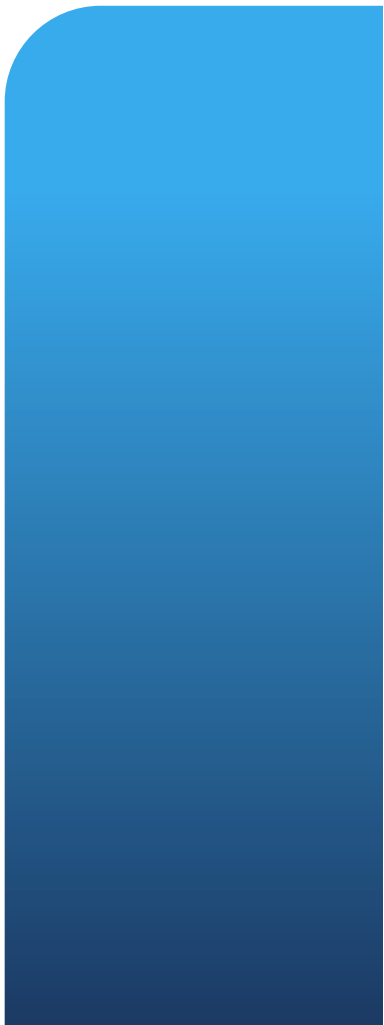


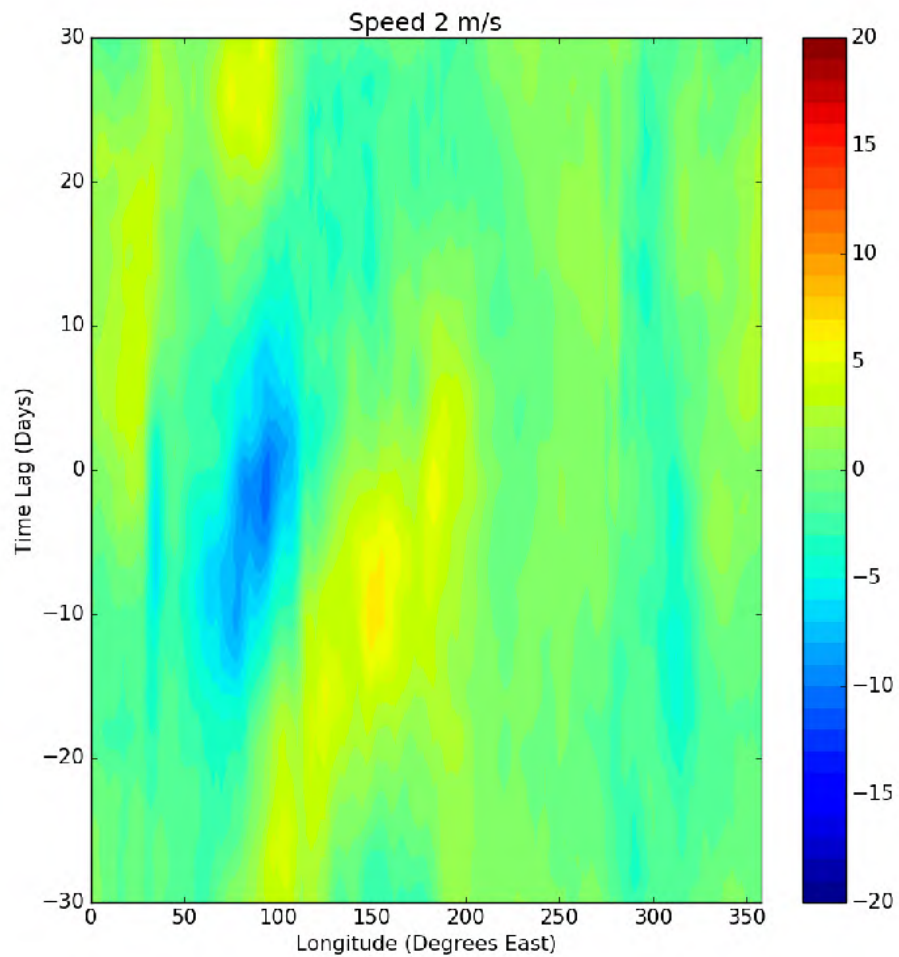
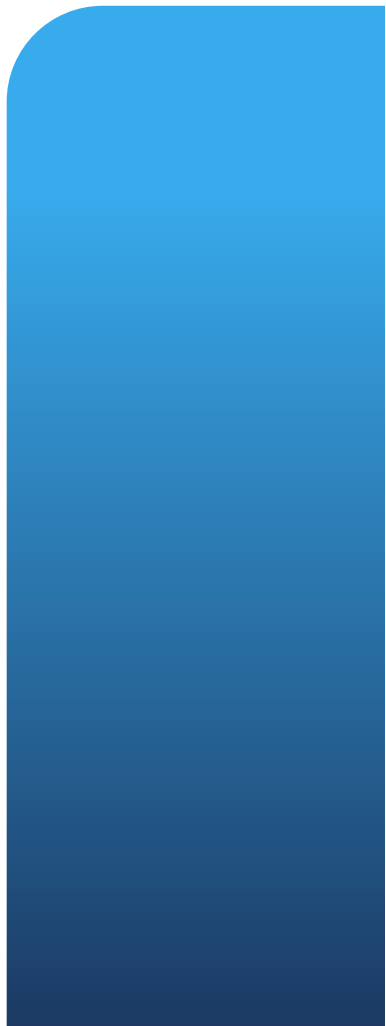
Conclusions

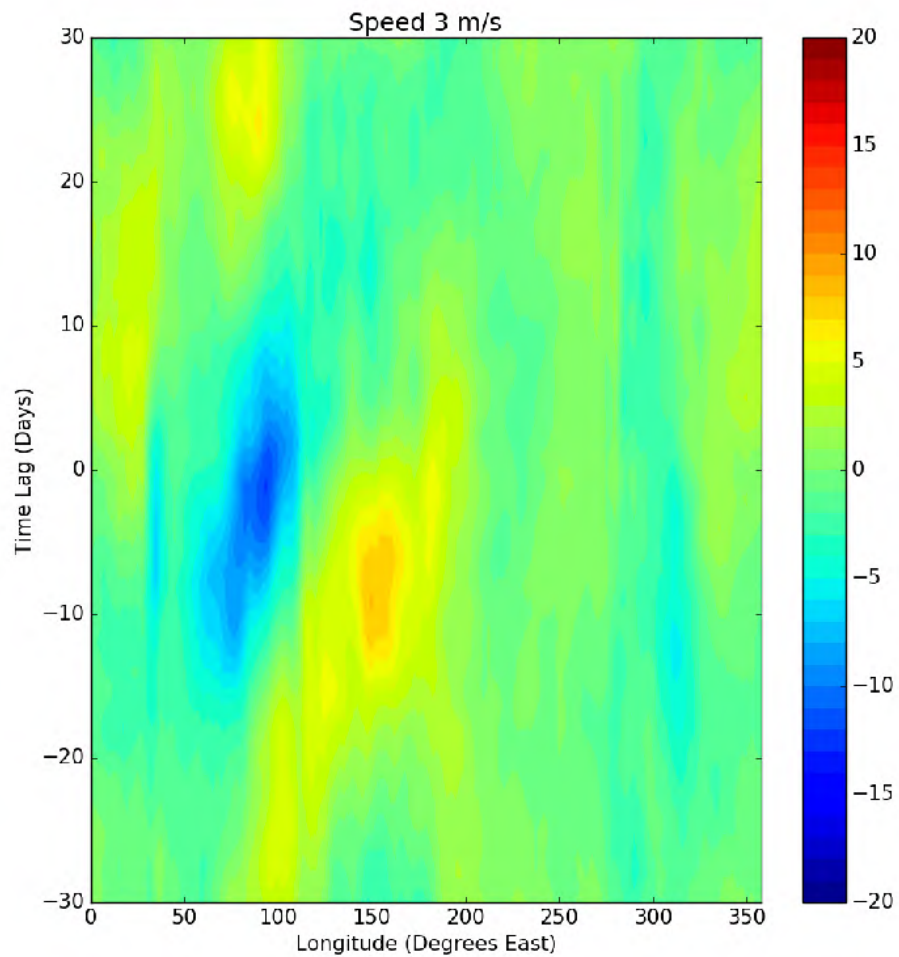
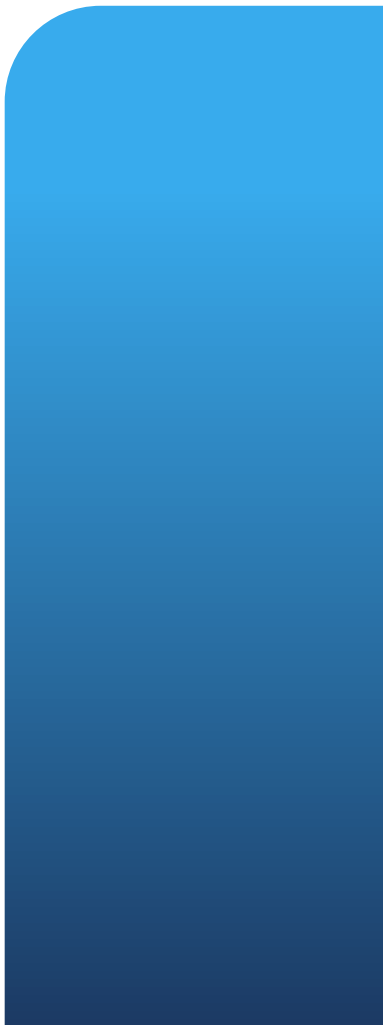
- MJO and Kelvin wave signals exist as parts of a continuum within a single broad spectral peak.
- MJO impact on extratropical circulation is well known
- Extratropical patterns are relevant to the propagation characteristics of the observed MJO: They are not just passive response patterns

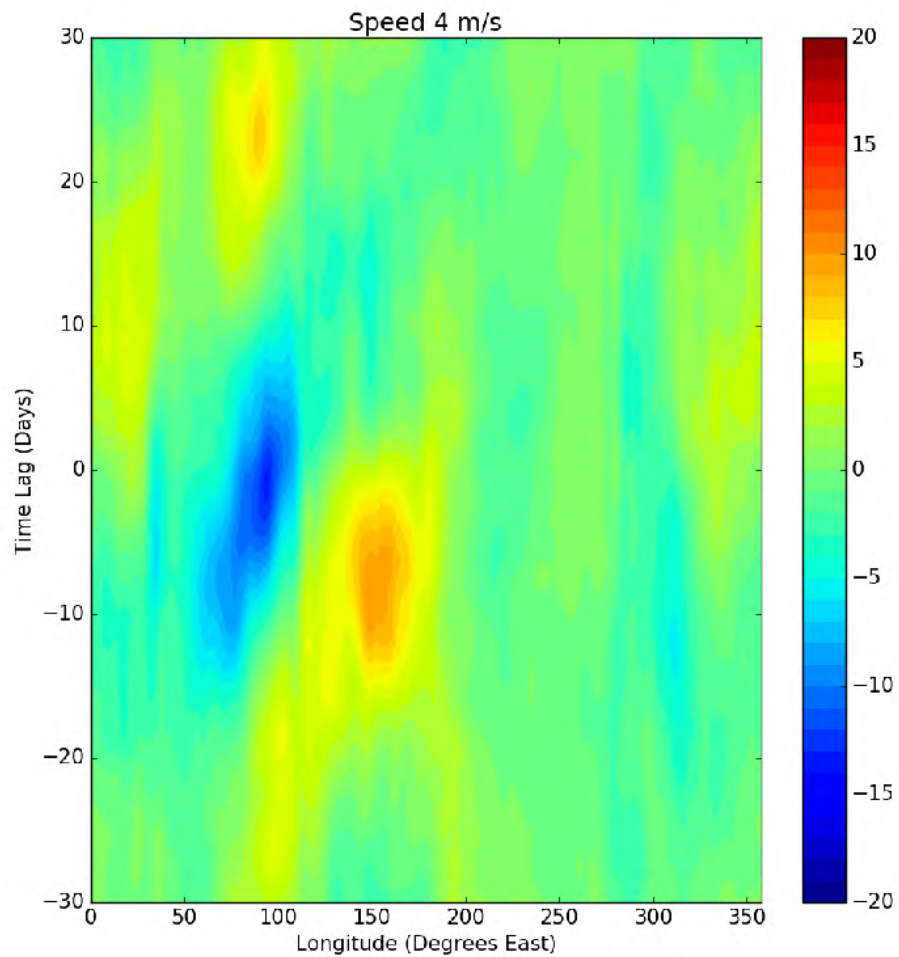
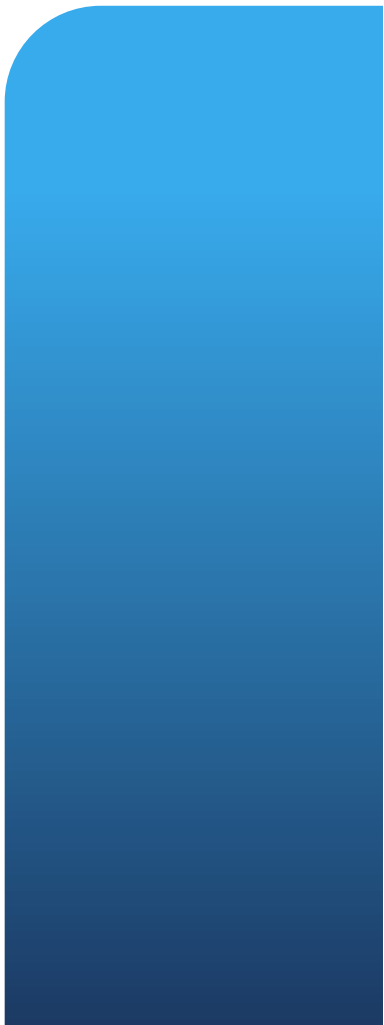
- The slowest MJO events over the Indian Ocean have cool dry air intrude over the Indian Ocean, facilitated by blocking over western Europe. Slow events have slow subtropical jet at proximate longitudes.
- The fastest Indian Ocean MJO events have more blocking over the Pacific Ocean and fast background subtropical jets at proximate latitudes.

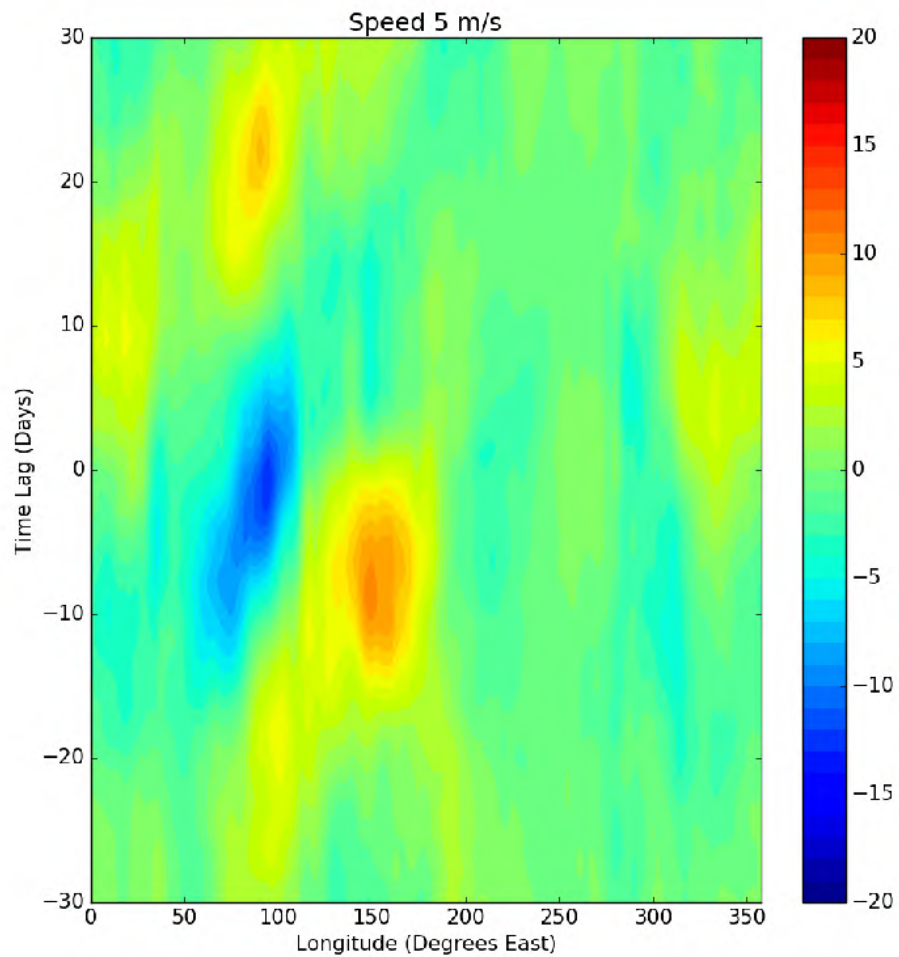
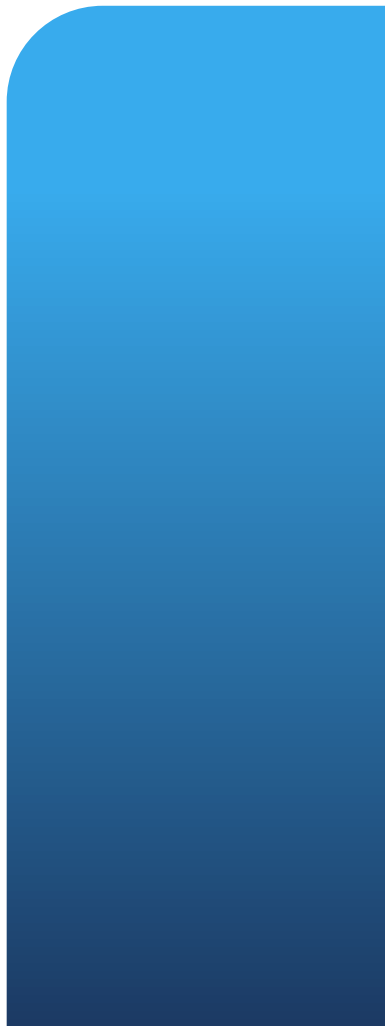
- The spectrum of observed intraseasonal variability must depend on the background state circulation.
- If models depend similarly, than getting the spectrum right must account for these effects. Concentration on convection is not enough.

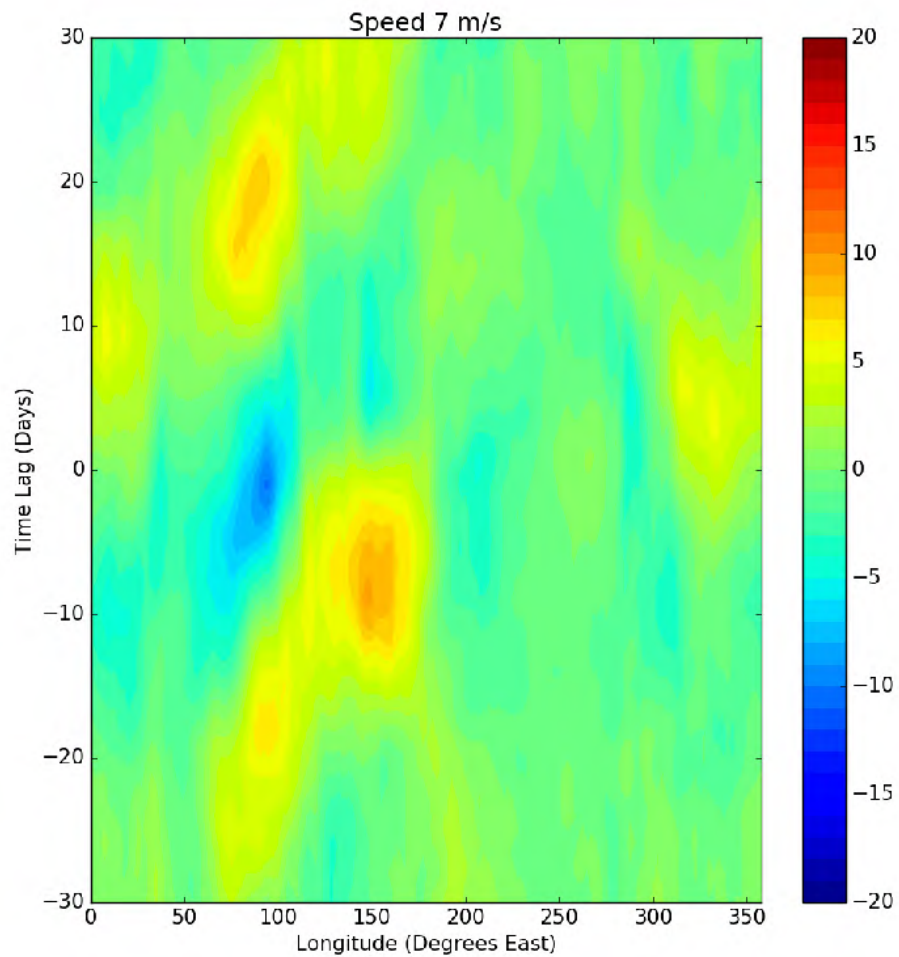
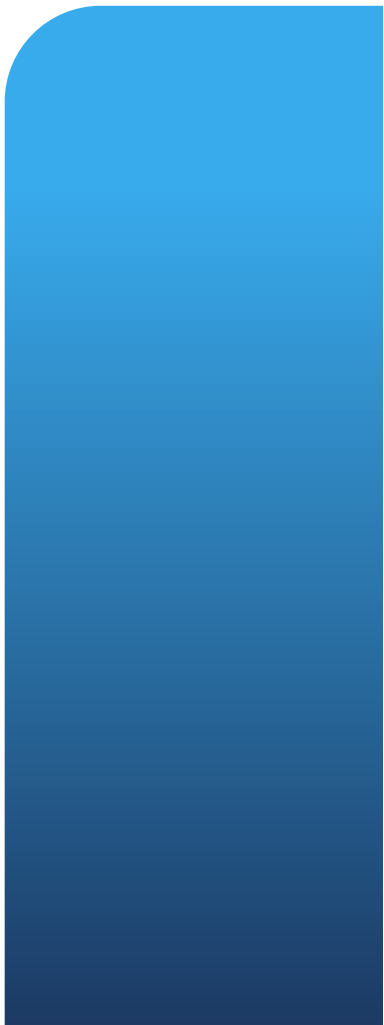


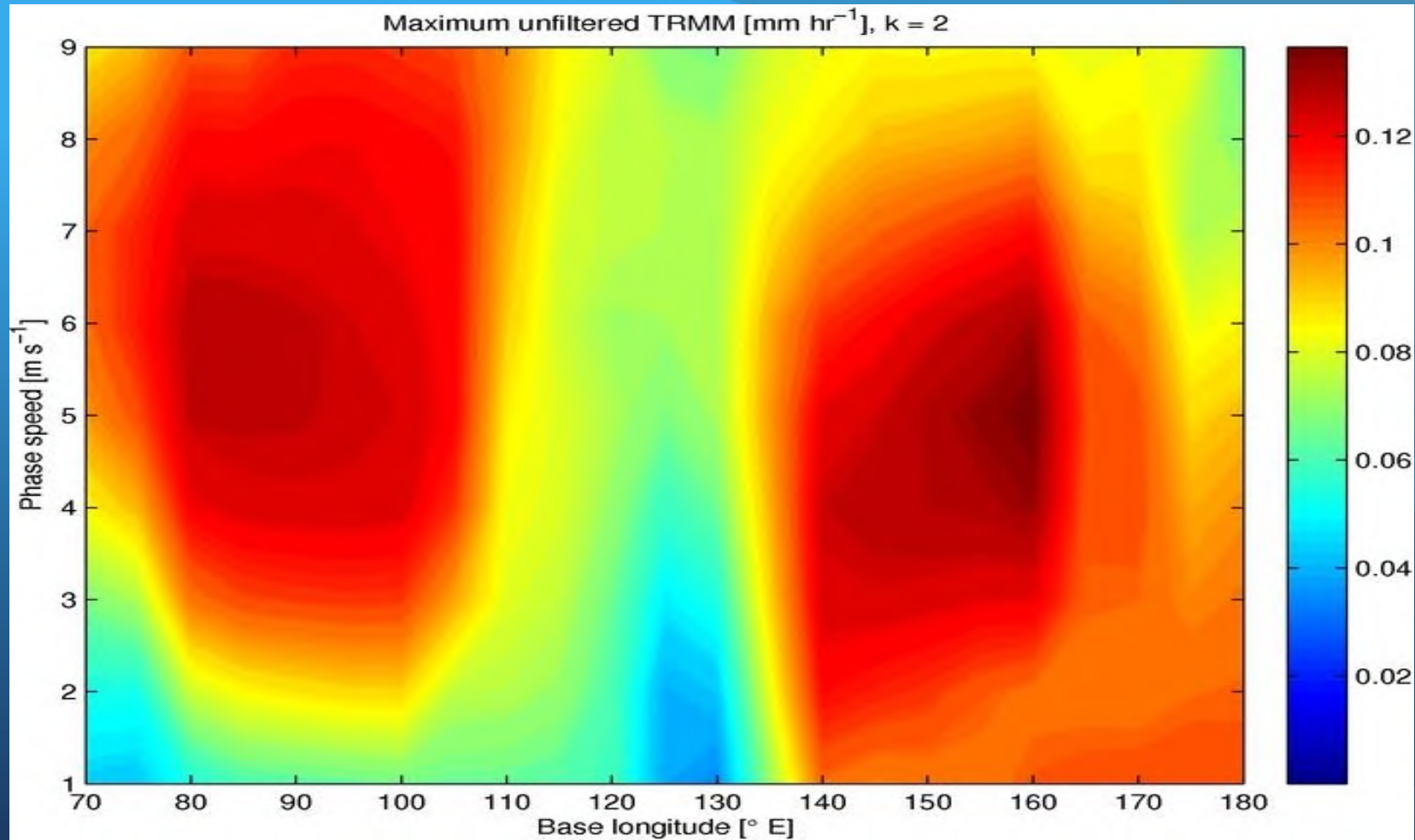




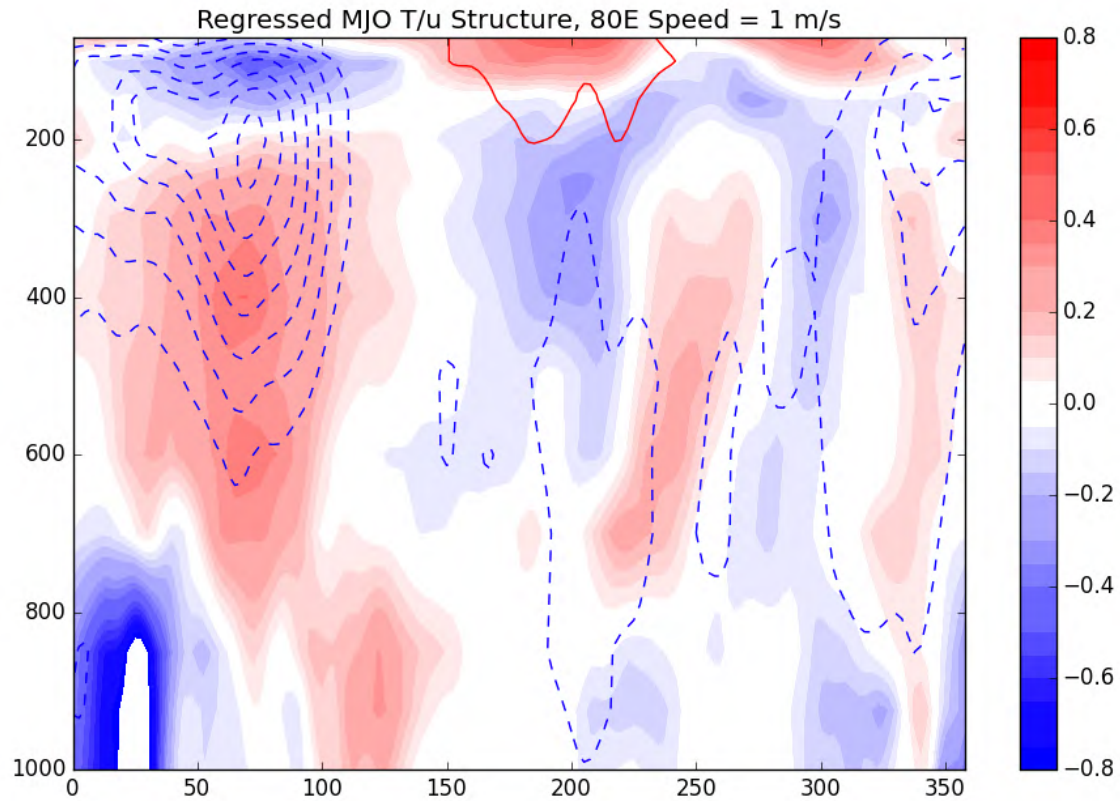




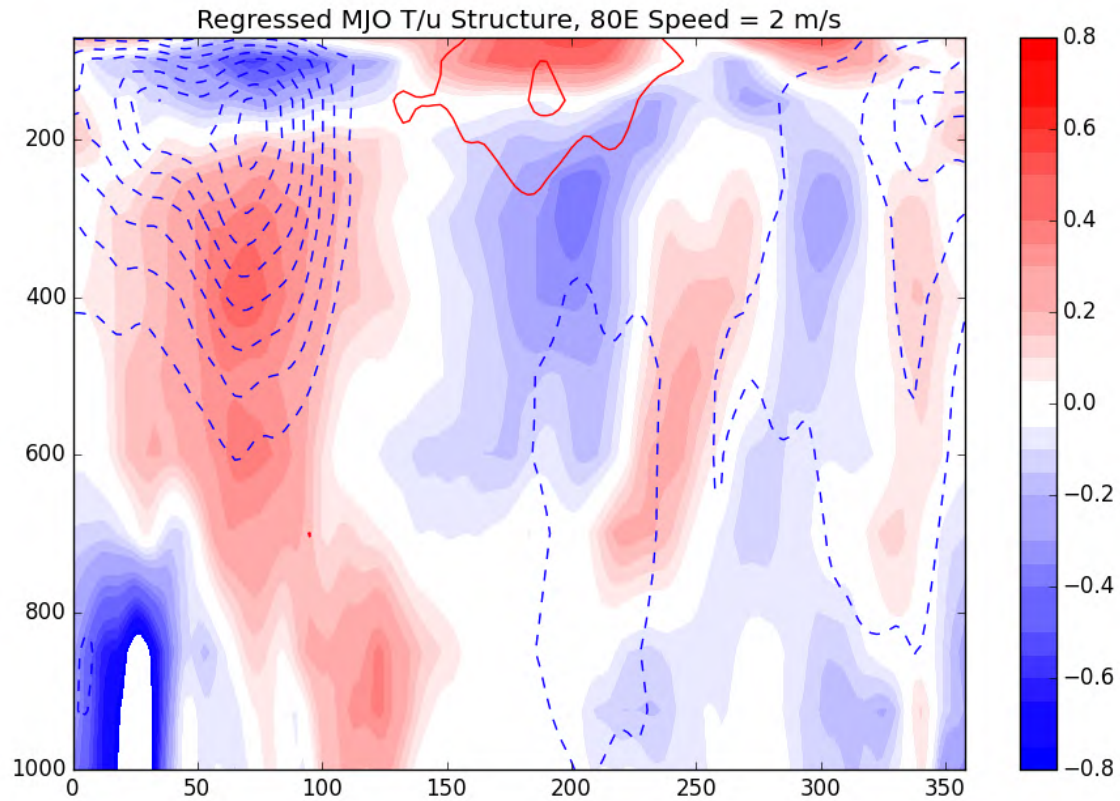




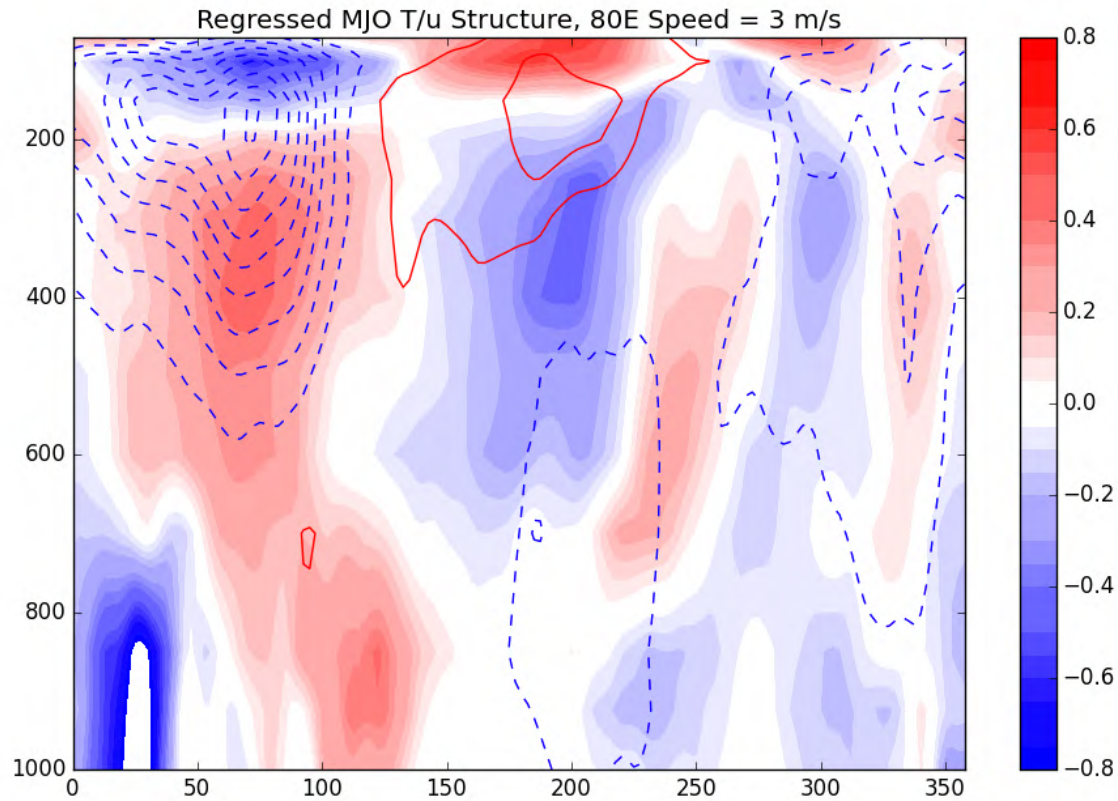
15° N



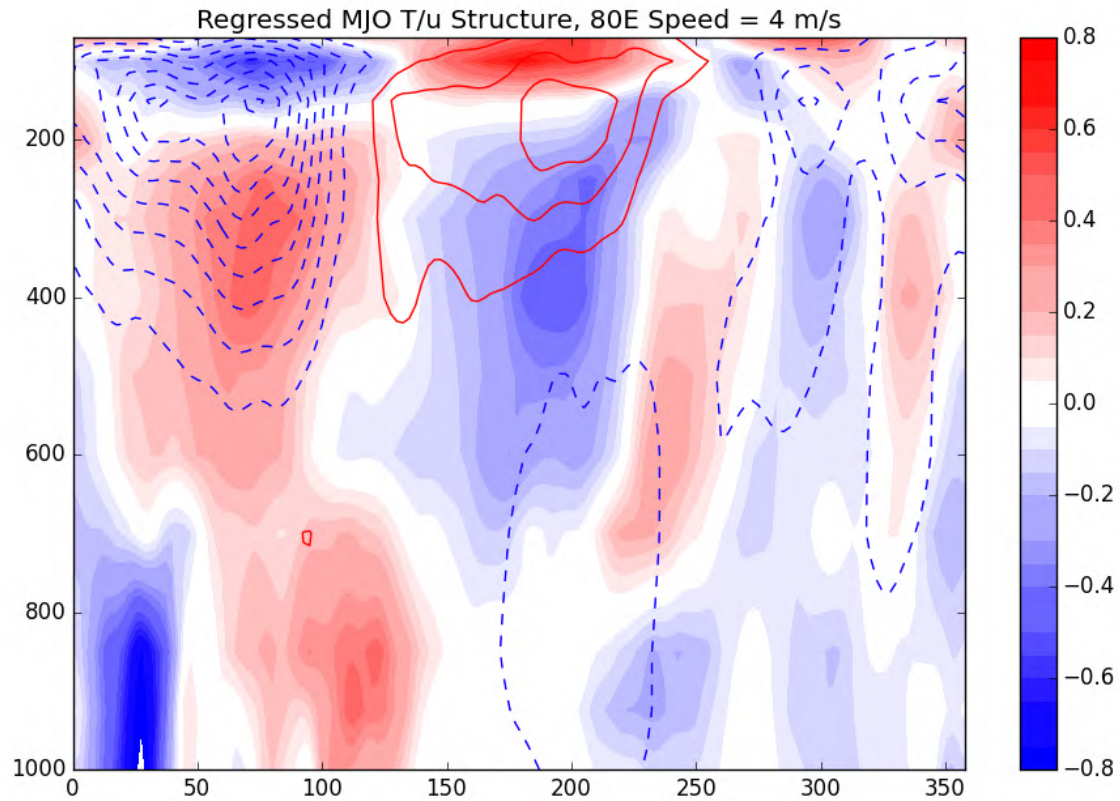
15° N



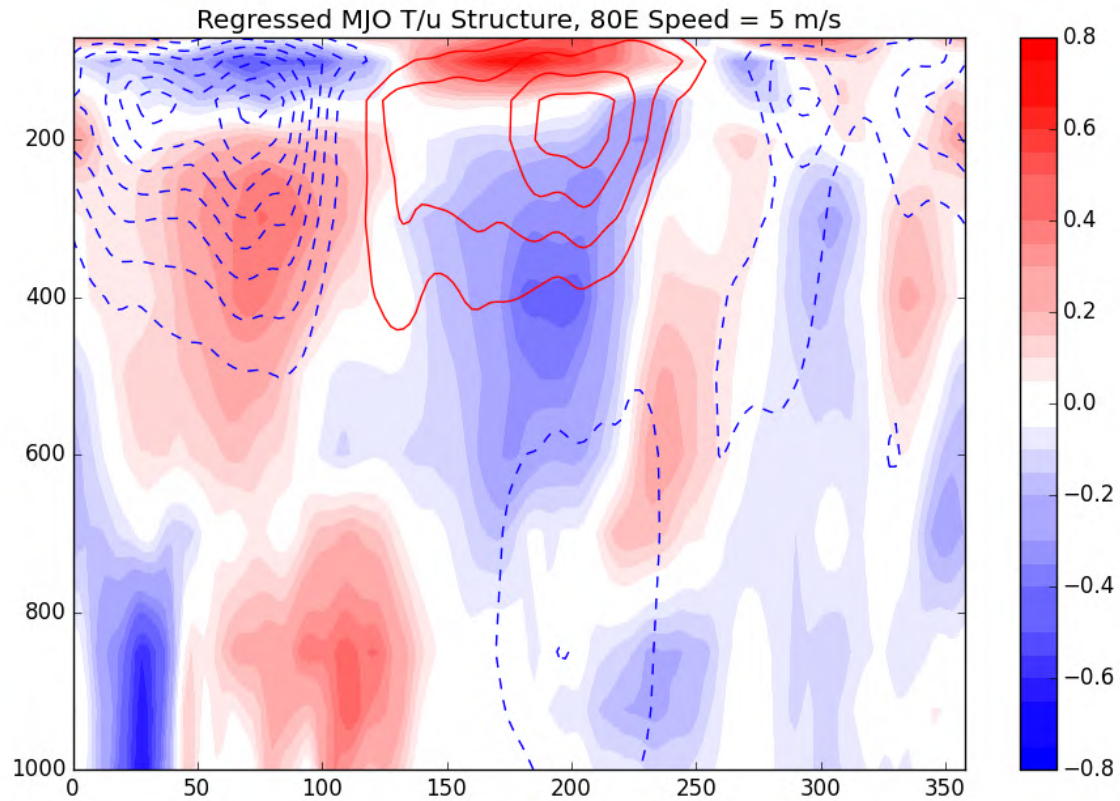
15° N



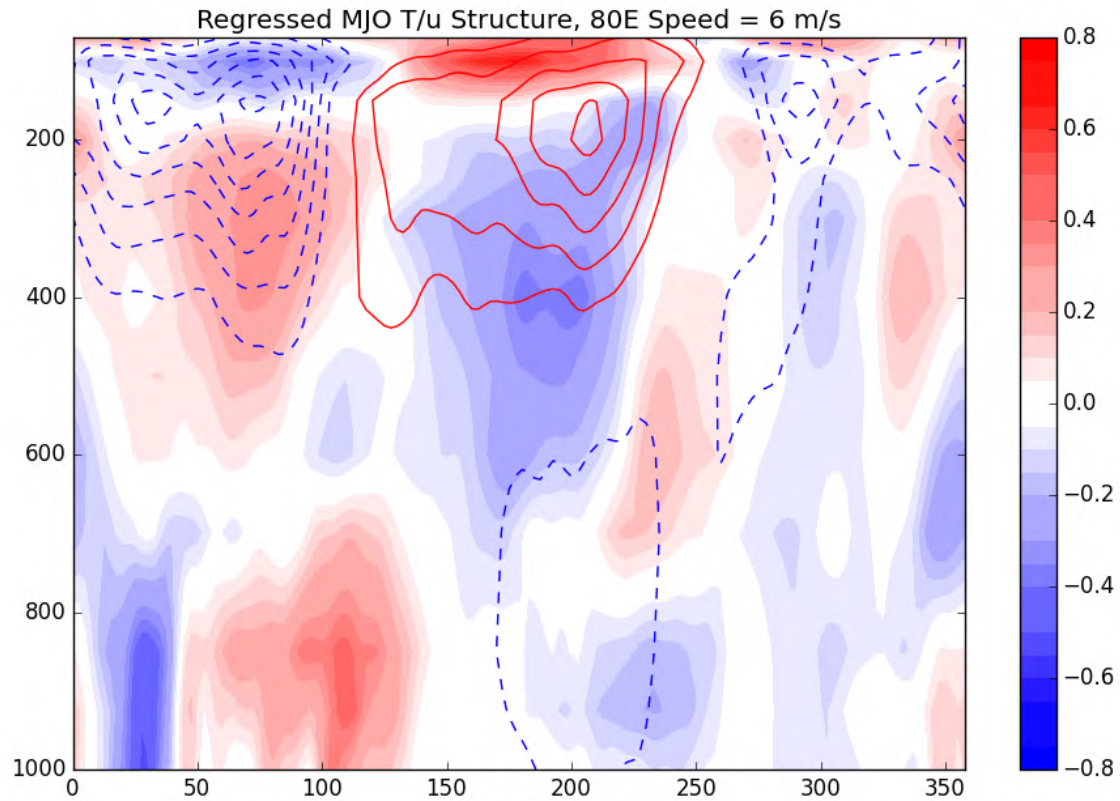
15° N



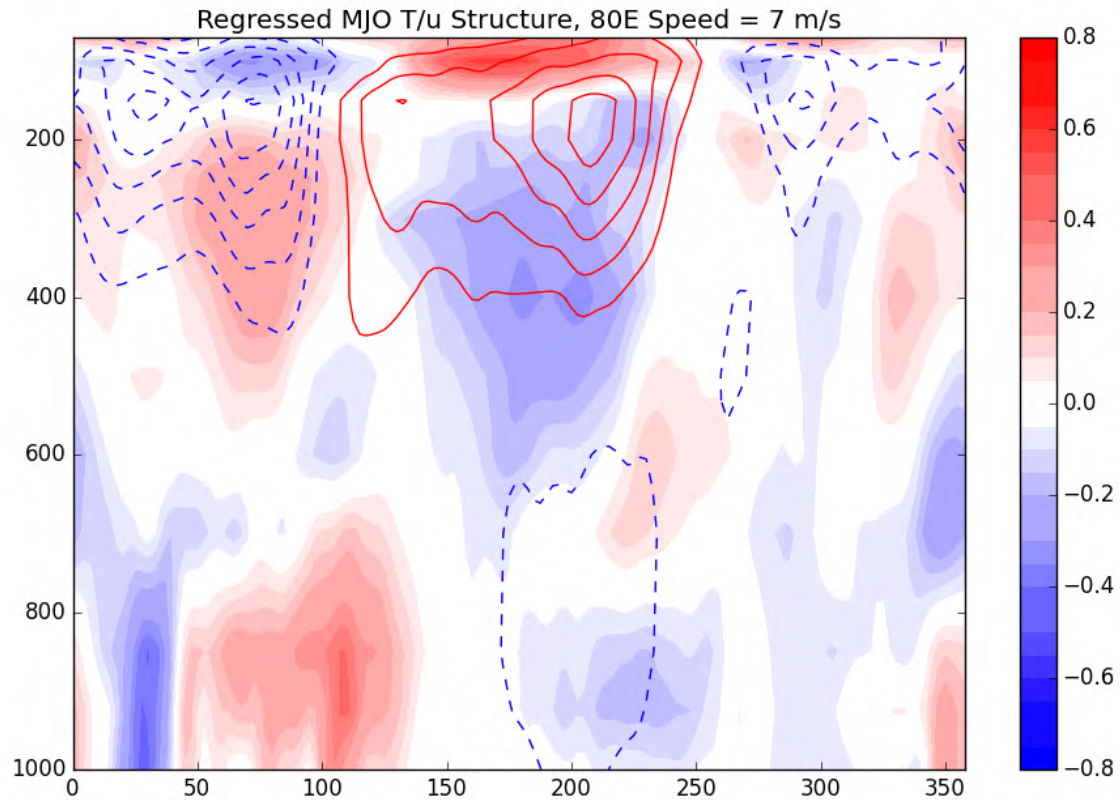
15° N



15° N



15° N



15° N

